

Spec	ial Instructions
1	-m360,sUN-11B,SOUTHEASTBRITISH.COLUBMIA
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4	IGEOLOGY., IN VICINITY, OF. THE. ORO. DENORO., MINE, (82E/2E) s10B +4,
5	rm420,By BN. Church
6	
7	im420,INTRODUCTION: M,Detailed mapping was initiated following renewed exploratio
8	and mining in the vicinity of the old Oro Denoro workings, 9.7 kilometres northeast
9	of Greenwood. Early in 1975 Granby Mining Corporation began a test operation to prov
0	an orebody estimated by previous work to contain about one million tonnes grading
11	slightly less than 1 per cent copper. When the property was visited in June 1975,
12	excavation had advanced to the third bench in the open pit, and approximately 135,00
13	tonnes of bedrock has been removed. Subsequently mining ceased pending the results o
14	a percussion drilling program. In 1976 stockpiled ore at the pit site was transporte
15	to the Phoenix mill.s+4pts,
16	iThe history of Oro Denoro can be traced to the original discovery of copper in the
17	so-called 'Summit Camp' in 1891. Beginning in 1903 the property became an important
18	local mine producing 136,447 tones of ore grading 1.37 per cent copper, 0.027 ounce
19	per ton gold, and 0.225 ounce per ton silver. By 1910 accessible ore had been extract
20	from five open stopes and 1,800 feet of underground drifts. After many years of
21	inactivity, prospecting was revived in response to increases in the price of copper.
22	Between 1951 and 1953, Attwood Copper Mines Ltd. carried out a number of geological,
23	geophysical, and geochemical surveys. Later the property was drilled by Noranda
24	Mines, Limited (1955 to 1957) and again by West Coast Resources Ltd. (1965 to 1970)
25	Testing by Granby Mining Corporation in the area continued through 1976.=10B,
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jPHYSIOGRAPHY:sM,,, The topography in the vicinity of the Oro Denoro mine is 2 relatively subdued, the hills in this region having been smoothed by southeasterly 3 moving Pleistocene glaciers. The mine site, at 3,500 feet elevation, occupies a low point on the ridge dividing the Boundary Creek and Granby River drainage systems. Slopes rise to the southwest in the direction of the Phoenix mine, attaining a 5 maximum elevation of 4,500 feet m.s.l. on Deadman Hill. Low points in the area near 6 Eholt Creek on the north, Lion Creek on the west, and Wilgress Lake on the northeast 7 are characteristically well vegetated and drift covered, bedrock being exposed mostly 8 9 in road cuts.s+4pts. 10 Access is from a number of old railway grades, now converted to logging roads, and 11 Highway 3 which bounds the map-area on the north and east. s10B, 12 13 'GENERAL, GEOLOGY: M,... The Oro Denoro region is underlain mostly by Mesozoic beds east of the drainage divide and an older basement complex of resistant metamorphic 14 (f_{U}^{\prime}) rocks to the west. These units are intruded by an assortment of plutonic rocks 15 16 ranging from granodiorite to gabbro, serpentine, and smaller dykes and sills of mostly diorite and pulaskite composition. +4pts, 17 18 'The relative age of formations is judged from metamorphism and the cutting relationship 19 of intrusions. Fossil evidence and a few radiometric determinations provide some 20 specific control. 101, 21 22 'METAMORPHIC, BASEMENT .COMPLEX: .M, ... The basement rocks, comprising an assemblage 23 of amphibolites, quartzies, marble bands, gneiss, and schist are collectively referred to as the 'Knob Hill Group.' The age range of the constituent formations 24 is unknown although they appear to be archaic and are certainly pre-Mesozoic. +4pts, 26 'Amphiboites predominate in the southwest part of the map-area. These are dark-coloured 27 and generally massive, medium to fine-grained rocks of probable basic volcanic 28 derivation. In thin section there is little evidence of primary textures or mineralogy, 29 these features having been largely obliterated by cataclasis and regional metamorphism. Fresh samples consist of fine-grained aggregates of plejochroic green **3**0 31 and yellow amphibole accompanied by thin discontinuous bands and lenses enriched in 32 magnetite dust or plagioclase. Retrograde effects have commonly reduced the rocks in schure to a **bash** of chlorite, carbonates, and clay minerals. Chemical analysis is the 33 main vestigial evidence proving the basaltic nature of the rocks (sI, seesM, 34 analysis No. 1 in the-accompanying-table).s+4pts,

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1	iThe most prominent outcrops of metaquartzite are near the southwest corner of the
2	map-area on the crest of the ridge leading to Deadman Hill. Characteristically the rock
3	is competent and ranges from dark grey to light cream coloured. In thin section
4	samples are commonly very fine-grained and cherty except for numerous veinlets of
5	coarser quartz.s+4pts,
6	The amphibolites and metaquartzite are accompanied by an older complex of gneiss
7	and schist. This basal assemblage is exposed on the hillside northeast of Lion Creek.
8	Generally the rocks are well foliated having thin alternating bands of felsic and
9	mafic minerals. In thin section a sample displays a mixture of small grains, averaging
0	about 0.1 millimetre in diameter, of quartz 25 per cent and biotite 40 per cent,
1	interspersed with composite granules of muscovite and quartz 35 per cent, and
2	accessory magnetite, s+4pts,
3	jThe fabric of the basement complex is often difficult to identify owing the massive
4	habit of the major units, such as the amphibolites, contortion of the gneiss and
15	schist formations, and paucity of marker beds.s10B,
16	1:4
17	jBEDDED,, ROCKS: sM,,,, Mesozoic and Tertiary strata great unconformably on the basement
18	complex. The Triassic Brooklyn Formation, lowermost in the cover assemblage, consists
19	mostly of limestone and clastic sedimentary rocks and is overlain by the slightly younger
20	Eholt Formation, & predominantly volcanic unit. The Tertiary beds, consisting of two-
21	formations, the Marron volcanic rocks and Kettle River sedimentary rocks, are of minor
22	importance, occurring only as small outliers.s101,
23	
24	jBROOKLYN,,FORMATION:sM,,,,Two members constitute the Brooklyn Formation,- at the
25	base, a widely distributed 'sharpstone conglomerate' and, uppermost, a thick limestone
26	deposit+4pts,
27	The sharpstone member forms a narrow belt trending subparallel to the crest of the
28	northeast ridge of Deadman Hill. It is a thickness of about 1,500 feet
29	of well indurated and often massive pebble conglomerate. Detailed examination of the
30	constituent fragments indicates a diverse provenance. Chert and greenstone are most
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		with			
1 abundant, compr	ising about 80 per cen	t of the clasts carbo	onate, schists, and gneiss		
2 fragments are a	ccessory. Modal analys	is of the sandy matrix	k shows an average of cher	t	
3 and quartz grain	ns, 40 per cent; amphi	bolite and porphyritic	c clasts, 25 per cent;		•
4 schist and chlo	futtul rite hash, 15 per cent	; carbonates, 15 per o	cent; and minor feldspar a	nd	
5 iron oxides. Ca	lculations based on ch	emical analysis of a	sample of carbonate-poor		
6 sharpstone cong	lomerate (analysis No.	2) gives 47 per cent	normative quartz.~+4pts,		
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1 iContact relations of the sharpstone member are displayed on the slopes of Deadman 2 Hill. At the lower contact the conglomerate directly overlies Knob Hill quartzites, 3 and at one point, a thin wedge of felsic tuff breccia. The upper contact, exposed 4 further east on the main ridge, passes transitionally into the limestone member 5 through several hundred feet of intercalated sandstone, conglomerate, and carbonate 6 beds. +4pts,

7 'The limestone member, estimated to be about 2,000 feet thick, is exposed extensively 8 in the eastern section of the map-area. This is a light blue-grey rock of variable 9 structure consisting of massive relatively pure calcium carbonate phases and thinly 10 bedded zones enriched in clay and chert impurities. Above the sharpstone contact 11 in the area southwest of Wilgress Lake and the Oro Denoro mine, the member is commonly *lind variables* 12 massive, in places having the aspect of a coarse breccia. To the east, the upper half 13 of the limestone section is generally well bedded with frequent shaly partings. Cherty 14 sand 'section of this peculiar sedimentary rock contains mostly rounded carbonate 16 clasts and fine-grained carbonate mud matrix, 55 per cent; subangular to well-rounded 17 chert grains, 35 per cent; and accessory quartz, feldspar, porphyritic rock fragments, 18 and amphibolite.st4pts,

19 'The age has been determined as Middle Triassic by the discovery ofsI, DaonellasM, sp. 20 in the limestone near the Phoenix mine (H. W. Little, personal communication).s10I, 21

22 iEHOLT..FORMATION:sM,...,The Eholt Formation, described by Carswell (1957) is little
23 known in the map-area. The only exposures are near the Phoenix road turnoff and rock
24 cuts along Highway 3. The formation consists of intercalated limestone and dark green
25 and purple lavas and breccia. A peculiar breccia facies of mixed chert, greenstone,
26 and limestone blocks (similar to the Brooklyn sharpstone conglomerate) has been
27 included in the definition of the formation. Chemical analysis of a sample of
28 Eholt lava gives a basaltic composition (analysis No. 3).st4pts, hrg.ke/
29 iThe Eholt rocks are believed to unconformably overlie Brooklyn Formation. The age
30 has beendetemind approbable Late Triassic form an assemblage of fossil corals,
31 which includessI, ThecosmiliasM, sp. (H.,W. Little, personal communication), obtained
32 from an outcrop of Eholt limestone north of the Phoenix road.s101,
33 s,

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	<u>1 'KETTLERIVE</u>	RFORMATION: SM,The b	asal Tertiary assemblag	e represented by the
	2 Kettle River	Formation is found only	in a small area north o	f Wilgress Lake. This
	3 is a channel	deposit of ligt-coloured	sandtone and pebble co	nglomerate. A thin
	4 section of th	e sandstone shows the fo) llowing modal compositi	on: quartz, 15 per cent;
	5 chert, 5 per	cent; feldspar, 40 per c	cent; volcanic rock, 10	per cent; mica and opaque
	6 minerals, acc	essory; matrix, 30 per c	ent. It is estimated th	at more than 75 per cent
^	7 of the clasts	were derived from a fre	esh felsic volcanic sour	ce rocks, the remainder
X	8 having a meta	Nort morphic, provenance.s101,	•	
[²	9	<i>[</i> .		
	10 IMARRON, FORM	ATION:sM,.,,The Marron v	volcanic rocks are found	l in two places, a small
	11 area of expos	ure on Deadman Hill and	north of Wilgress Lake.	At the latter locality
	12 Marron rocks	directly overlie the Ket	ttle River beds.s+4pts,	
	13 iThe volcanic	rocks are medium to dan	rk grey lavas and brecci	a characterized by
	14 scattered tab	ular or rhomb-shaped and	orthoclase and anhedral	pyroxene phenocrysts.
	15 In thin secti	on the matrix is common!	ly charged with randomly	v arranged feldspar
	16 microlites, r	ounded analcite crytals	, and interstitial pyrox	æne, magnetite, abundant
	17 apatite, and	glass. Chemical analysis	s of the rock demonstrat	es a mafic phonolite
		sI,seesM, analysis No.		
	19	Ta	inter	in the area,
	20 jICNEOUS,, INI	RUSIONS:sM,,,,The Lion	Creek granodiorite, the	main intrusion, is
		y smaller satellitic bo		
		Ofter intracens a	nglude	Coryell-related monzonite
	23 sills and pul	.askite dykes, c omplete t	he spectrum of tyneous	intrusions.s10I,
	24			
	25 jultramafic,	ROCKS:sM,,,,A small len	s of serpentinized perio	dotite occurring in basemen
	26 gneisses nort	th of Lion Creek constit	utes the only ultramafic	c body and probable oldest
	27 intrusion in	the map-area. These roc	ks are generally brittle	e and mottled light grey -
	 28 dark greenist	n grey on fresh surfaces	and rust brown where we	eathered. In thin section,
	29 typical samp	is le comprises a cataclast	ic aggregation of serpe	ntine, talc, and minor
		-	and	iated with magnetite. 101,
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	33 N	y i	A as dalla	on page 8.
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*****	2	 'LION, CREEK GRANODIORITE: ~M,, The Lion Creek granodiorite is thought to be an
	3	appendage of the Wallace Creek batholith. The intrusion enters the map-eres-from the
	4	est penetrating the Triassic beds and basement complex and extending eastward to the
	5	Oro Denoro workings. This is a light grey massive granoblastic rock having the
	6	following approximate modal composition: feldspar, about 65 per cent (mostly
	7	plagioclase); quartz, 25 per cent; and accessory amphibole, biotite, and magnetite.
	8	The plagioclase occurs as rectangular oscillatory zoned plates, intermixed with
	9	subhedral quartz, 1 to 4 millimetres across, with slightly smaller interstitial quartz
	10	and feldspar and magnetite grains associated with patches of pale yellow-green
	11	amphibole and scattered biotite books. Some garnet, sphene, and diopside are reported
	12	near the contaminated borders of the intrusion.s+4pts,
	13	The Emma intrusion is an elongated northerly trending offshoot of the Lion Creek
	14	body. It is well exposed in a small area west of the Oro Denoro pit and on the
	15	slopes to the north in the vicinity of the Emma mine. This is a distinctive
	16	porphyritic phase consisting of about 8 per cent subhedral plagioclase individual
	17	and clusters, and 1 per cent partly resorbed amphibole phenocrysts, $\frac{1}{2}$ to 4
	18	millimetres in length, suspended in a fine-grained quartzofeldspathic matrix.
	19	According to normative calculations the rock contains a total of 20 per cent quartz
	20	and about 70 per cent felds apr (Ansdown, 30 sup,), the chemical composition of the
	21	Emma phase being very similar to the Lion Creek intrusion (sI,seesM, analyses
	22	Nos. 6 and 7).s+4pts,
	23	Nos. 6 and 7).s+4pts, x $-\pi lle l$
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1	iThe age of the Lion Creek intrusion has been determined as Cretaceous,
2	140smath,bsM,5 Ma, based on K/Ar analysis of biotite obtained from a sample of
3	granodiorite taken near the Emma mine.sI,
4	
5	iCYCLOPS,.GABBRO: M,,, The Cyclops gabbro occurs in the area south and southeast of the
6	Cyclops prospect as a large partly concordant body and a number of small dyke-like
7	offshoots in the vicinity of the Oro Denoro and Emma workings. The rock is commonly
8	dark greenish grey and rather uniform, fine grained, consisting of subhedral plagioclase
9	plates, about 55 per cent, interspersed with equant pyroxene grains, 20 per cent
10	(measuring to 1.5 millimetres in diameter) set in a matrix of chlorite and disseminated
11	magnetite. Conversion of some of the pyroxene to blue-green amphibole locally is
12	viewed as a retrograde metamorphic effect.s+4pts,
13	The age of the gabbro certainly post-dates the Middle Triassic Brooklyn Formation which
14	it intrudes, however, its relation to other igneous rocks in the area is uncertain.s10I
15	,
16	iCORYELL,, INTRUSIONS: SM,,,. The Coryell intrusions are exposed mostly in the northern
17	part of the map-area. These include an assemblage of syenite, monzonite, and shonkiniti
18	bodies, and their finer grained equivalents,- a variety of pulaskite and lamprophyre
19	dykes. The most common rock in this suite is a mottled pink and grey feldspar $$ $$ $$
20	porphyry consisting of glomerophenocrystic plagioclase-sanidine clots, measuring-to-
21	6 millimetres in diameter, and smaller solitary feldspar crystals suspended in a
22	finer grained matrix of interlocking feldspars and biotite, and a small amount of
23	interstitial quartz, disseminated magnetite, and apatite. Clinopyroxene is also
24	present as an additional mineral in the more basic phases.s+4pts,
25	jA sample of pyroxene monzodiorite obtained northwest of the Emma mine shows the
26	following normative mineralogy: quartz, 6.5 per cent; potassium feldspar, 27.2 per
27	cent; plagioclase (Anodown,33 p , 41.1 per cent; clinopyroxene, 22.8 per cent;
28	magnetite, 2.4 per cent. The chemical analysis of this rock is similar to the compositi
29	of a post-mineralization pulaskite dyke exposed in the No. 1 quarry at the Oro Denoro
30	mine («I,seecM, analyses Nos. 8 and 9).c+4pts,
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	(1 173)
1	The age of Coryell batholith, recently established by James A. Fyles in the Rossland
2	area, is Middle Eocene.«10B,
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4	'STRUCTURE: M, The general pattern of folding, faulting, and intrusion seems to have
5	a north-south and east-west control. In detail, the structures are intricate and often
6	difficult to unravel because of the scarcity of markers and imperfect exosure.e+4pts,
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1	'The trend of the Brooklyn and Eholt Formations is mostly northerly, bedding
2	strikes averaging 005 degrees. The Mesozoic section is tilted easterly about 50
3	degrees forming a monocline. Local reversals and deflections of beds give evidence
4	of the presence of northeasterly plunging minor folds (Fig. 2).s+4pts,
5	iThe structure of the basement complex is more difficult to determine because of
6	the massive character of the rocks, particularly the amphibolite formation which
7	covers a wide area. An exception is a band of marble, traceable for about 2 kilometre
8	striking from the west boundary of the map-area to the north slope of Deadman Hill,
9	Und following near the base of the amphibolite formation. Metaquartzite with an average
10	bedding attitude ofschange to degree ball, 060!/50!schange 10M, northwest overlies
11	the amphibolite near the crest of Deadman Hill.s+4pts,
12	iThe main fractures trend easterly coincident with major draws and valleys, and
13	northerly subparallel to the principal strike direction of the Mesozoic beds (Fig. 1
14	Specific measurements of numerous minor fractures illustrates this bimodal
15	distribution showing a development of strong joint sets at approximatelysdegree ball
16	110!/85! southwest and 005!/82!schange ball, east (Fig. 3). That the fracture system
17	pre-dates the major igneous intrusions and was probably advantageous to the emplacem
18	of these rocks is suggested by the east-west elongation of the Lion Creek pluton
19	and the north-south orientation of the Emma and Cyclops bodies.s10B,
20	within
21	jMINERALIZATION: SM,,,, The Oro Denoro mine is centrally located on a 2.4-kilometre-lo
22	moridianal alignment of skarn deposits which includes the Emma and Jumbo on the nort
23	and the Cyclops and Lancashire Lass on the south. The host rock for all of these,
24	and many other deposits in the Greenwood area, is the Brooklyn limestone member
25	(Fig. 1).s+4pts,
26	jThe geology of the Oro Denoro is relatively straightforward. Mineralization consist
27	of pockets of pyrite, chalcopyrite, and magnetite in a garnetite skarn. This is most
28	a replacement of limestone intruded by an apophysis of the Lion Creek granodiorite
29	stock (Fig. 4).~+4pts,
30	'The mine workings cover an area of about 10 acres in the central part of the Oro
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1 Denoro Crown-granted claim. In the early period of mining between 1903 and 1910 2 ore was drawn from five quarries and a number of open stopers which were serviced by 3 from two underground levels. This, and the Phoenix operation several miles to the

southwest were among the earliest attempts at open-pit mining in the Province. +4pts,

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1 'In the old workings of Oro Denoro the southernmost quarries, Nos. 1 and 2, were 2 the principal source of copper ore. These are interconnected and have a general 3 east-west elongation. The trend of the excavations appears to follow the course aling 4 of a number of large steeply dipping calcite lenses in the skarn and the granodiorite 5 contact which is near the north wall. Quarry No. 3, centred about 200 feet north of 6 Nos. 1 and 2, is the second largest pit. Here the mineralization was concentrated 7 in a tongue of skarn rock projecting deep into the granodiorite mass. Quarries 8 Nos. 4 and 5, centred about 150 feet northwest of No. 3, are relatively small. The 9 magnetite-rich ore was situated between a small remnant of limestone in the skarn rocks 10 and the granodiorite contract. Control of the mineralization appears to be east-west 11 crossfractures cutting approximately perpendicular to bedding in the limestone.s+4pts, 12 iThe new excavation which is located immediately west and south of the old quarries 13 is a large benched open pit about 500 feet long and 150 feet wide carved mostly 14 from garnetite skarn, forming the core and summit of Oro Denoro's 'mine hill.' The target of this development was a mineralized zone near the south end of the pit.s+4pts area iThe mine is traversed by a numer of faults of ore control significance. The 16 17 most important is a pronounced shear zone striking about 120 degrees from the north 18 end of the main pit through No. 1 quarry. Profound movement on this zone has resulted 19 in the emplacement of exotic formations in the skarn such as a wedge of carbonaceous 20 schist in the main pit and epidotized volcanic breccia along the south wall of No. 1 21 quarry. Of less importance are two minor faults 045/75 southeast and 015/80 east 22 causing local displacements in the skarn-granodiorite contact in No. 3 quarry and 23 a weaker fracture 150/45 southwest observed in No. 5 quarry.s+4pts, 24 jMineralization and the development of a skarn at Oro Denoro is evidently the result 25 of intrusion of the Lion Creek granodiorite stock. An exchange of chemical components 26 between the granodiorite and Brooklyn limestone is apparent. A determination of the 27 mineralogy of the skarn is provided by Carswell (1957): garnet 28 (grossularitesmath down,10sM up, andraditesmath down,90sM up), 35 per cent; 29 and 5 per cent for each of clinozoisite, diopside, and quartz (in weight per cent). 30 In terms of estimated chemical composition this mineralogy reduces to: 31 32 33 34

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1 S10smath down, 2⁻M up,. 39.6 per cent; Al-math down, 2 M up, 0⁻math down, 3⁻M up,,

2 3.9 per cent; Fe-math down, 2-M up, 0-math down, 3-M up,, 24.7 per cent; MgO, 0.9 per 3 cent; CaO, 30.9 per cent on an anhydrous base, - a calculation which compares closely with 4 the actual chemistry of a sample of skarn rock (analysis No. 4). The gain of large amounts -f dde f5 of iron oxide and silica by the limestone is matched by an equally large loss of lime to 6 the granodiorite. Source of the iron oxide and silica appears to result from 7 calcification of iron-bearing silicates and plagioclase feldspar in the granodiorite. 44, ÷ 8

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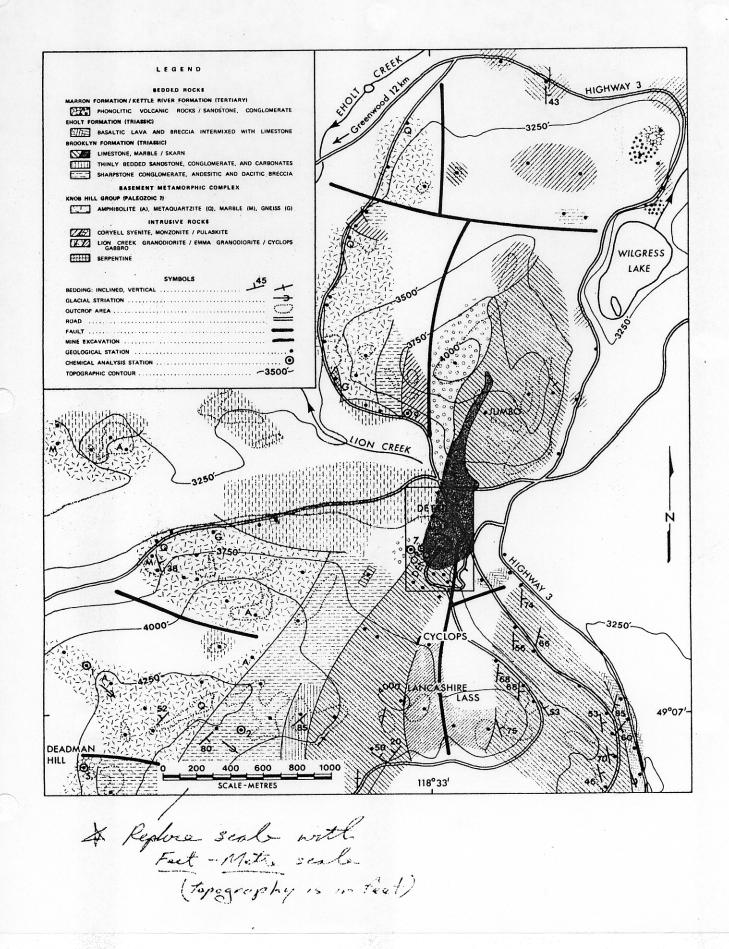
1 jm420, sUN-10M, Mineral paragenesis begins with the skarn silicates which are 2 partly overlapped by magnetite and succeeded by sulphides. Magnetite commonly occurs interbanded with the skarn silicates, the banding generally having the same trend 3 as bedding in the nearby limestone (Plate II). In contrast, the sulphides are 4 present as disseminated grains, individual large crystals or masses associated with 5 calcite in seams and pods (Plate III). Also, mixtures of sulphides, mostly pyrite 6 and chalcopyrite, with garnet and calcite may display a rude planar fabric or banding 7 of coarse and fine grains betraying a suggestion of relict bedding (Plate IV). The 8 final generation of sulphides is reposed in interstices and cracks crossing the 9 skarn silicates, magnetite, and older sulphides (Plate V).s+4pts, 10 includ iFormation of the skarn and emplacement of the sulphide and magnetite ores reform 11 12 to high temperature interaction between the limestone and granodiorite, - high 13 temperatures being implied by the extensive development of garnetite. Marked irregularity and variation in the width of the skarn zone from a few metres to many 14 suggests tens of metres derives the conclusion that the reactions occurred only at places 15 16 where ascending solutions were active. These solutions, enriched in carbon dioxide, 17 silica, and iron, rose along the fissure system. At first the interaction of 18 invading solutions on the host rocks was intensified along fissures converting with Cater atron of walls forming metasomatic veins, and thence infiltrating the limestone mass soaking 19 20 rock pores and ultimately achieving wholesale ceplacement. «10B, 21 22 IREFERENCES = 10M. 23 24 Assessment Reports 67, 117, 118.s+2pts, 25 oBrock, R..W. (1902): 26 Preliminary Report on the Boundary Creek District, #I, Geol. Surv., 27 Canada, M, Sum. Rept. +2pt, 28 •Carswell, H..T. (1975): 29 The Geology and Ore Deposits of the Summit Camp, Boundary 30 District, British Columbia, M.Sc. Thesis, sI, University of 31 32 33

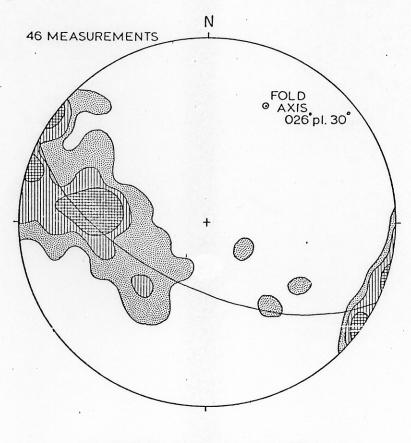
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Spec		
1		British Columbia, sM, 80 pp.s+2pts,
2	oDaly, R.,A. (191	2):
3		Geology of the North American Cordillera at the Forty-Ninth
4		Parallel,sI, Geol. Surv., Canada,sM, Mem. 38.s+2pts,
5	oFyles, James T.,	
6		Harakal, JE., and White, WH. (1973): The Age of Sulfide
7		Mineralization at Rossland, British Columbia,sI, Econ. Geol.,sM,
8		Vol. 68, pp. 23-33.s+2pts,
9	oLeRoy, 0.,E. (19	12): Geology and Ore Deposits of Phoenix, British Columbia, sI,
10		Geol. Surv., Canada, sM, Mem. 21.s+2pts,
11	oLittle, Hl.W. an	ad '
12		Thorpe, R.,E. (1965): Greenwood (East Half), sI, Geol. Surv., Canada, sM,
13		Paper 65-1.s+2pts,
14	oMcNaughton, D.,A	A. (1945): Greenwood-Phoenix Area,sI, Geol. Surv., Canada,sM,
15	• • •	Paper 45-20.s101 and +2pts,
16	oMinister of Mine	es, B.C., 5M,
17		Ann. Repts., 1965, pp. 171, 172; 1967, pp. 232, 233; 1968, pp.
18		233-235.s+2pts,
19		
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20		Forks-Eholt Area, Boundary District, British Columbia, M.Sc. Thesis, sI,
21		University of British Columbia, SM, 114 pp.s+2pts,
22		•
23		Geology and Copper Deposits of the Boundary District, British
24		Columbia, sI, CIM, sM, Bull., Vol. 49, No. 3, p. 684.s10B,
25		Columbia, Si, Cin, Sh, Bulli, 101. 47, No. 3, p. 004.3105,
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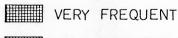


FIG. CHURCH - 4 DETAILED GEOLOGY IN VICINITY OF THE ORO DENO!





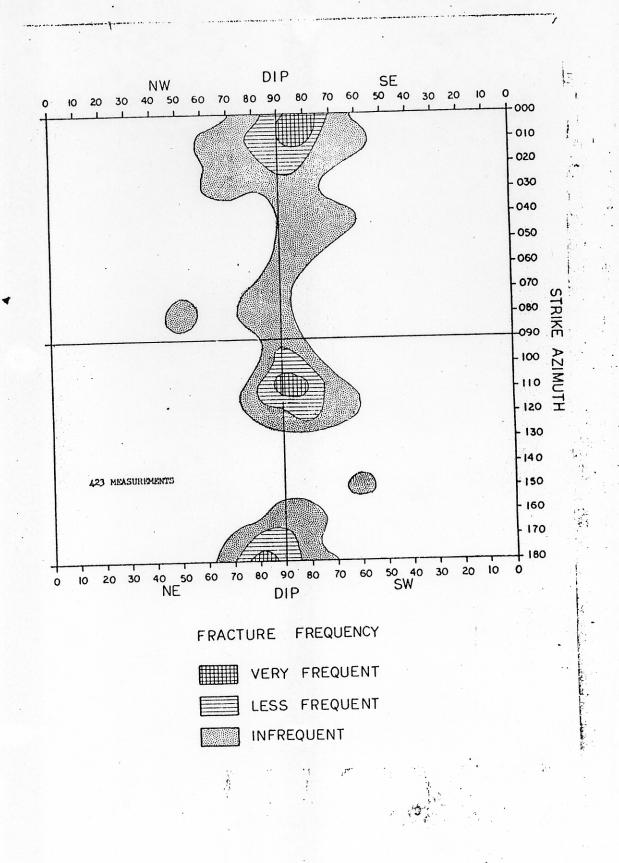
FREQUENCY OF POLES TO BEDS



LESS FREQUENT

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FIG. CHURCH - 2 EQUAL AREA PLOT OF BEDDING FOR MESOZOIC ROCKS, ORO DENORO AREA



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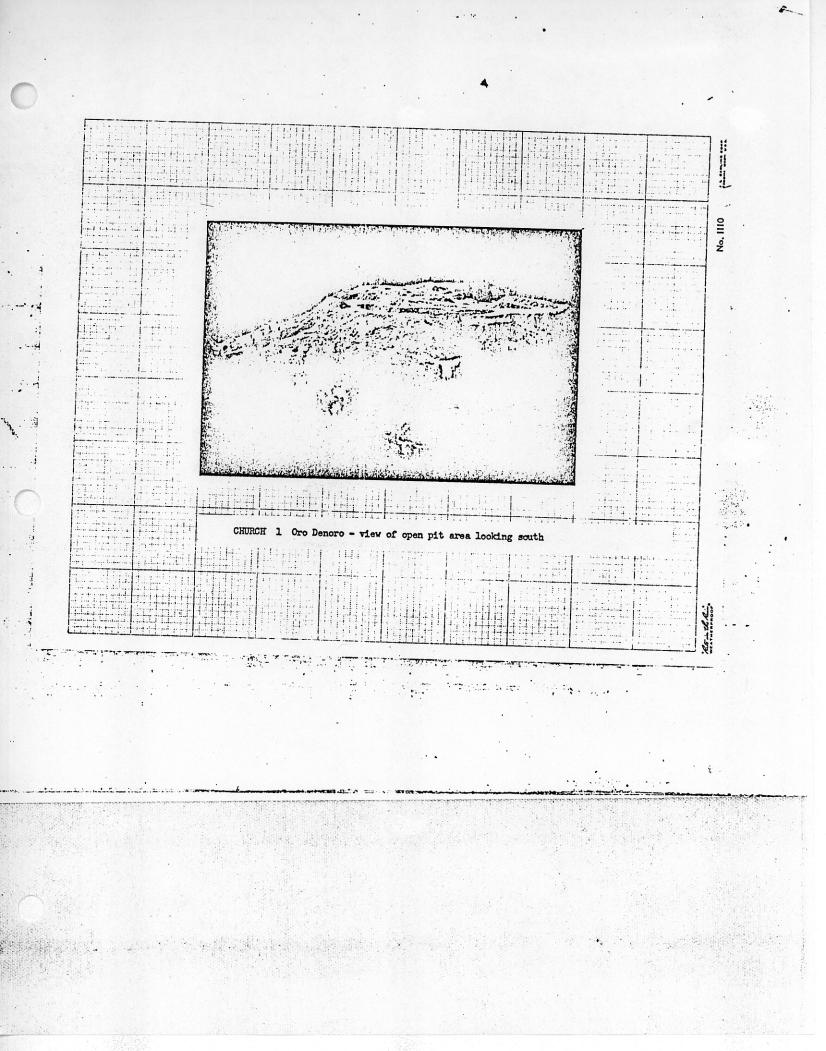
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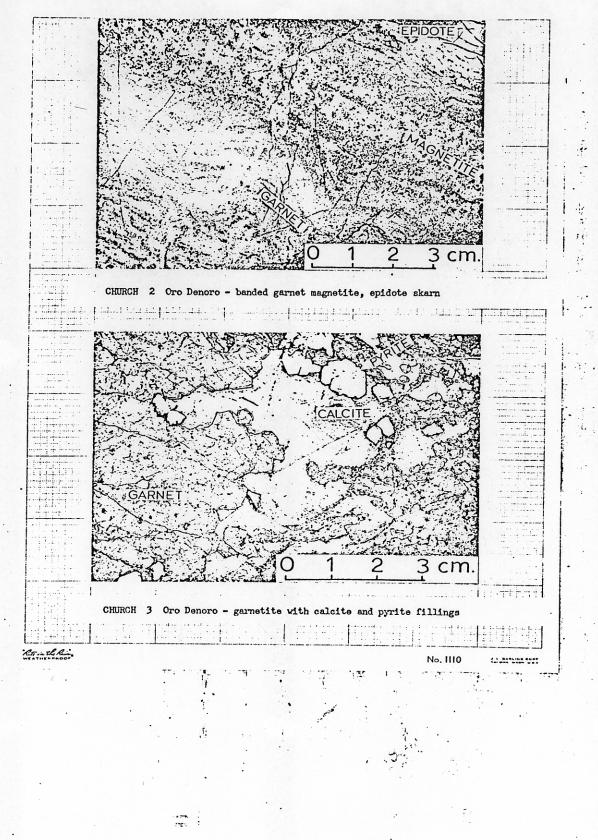
PLATES

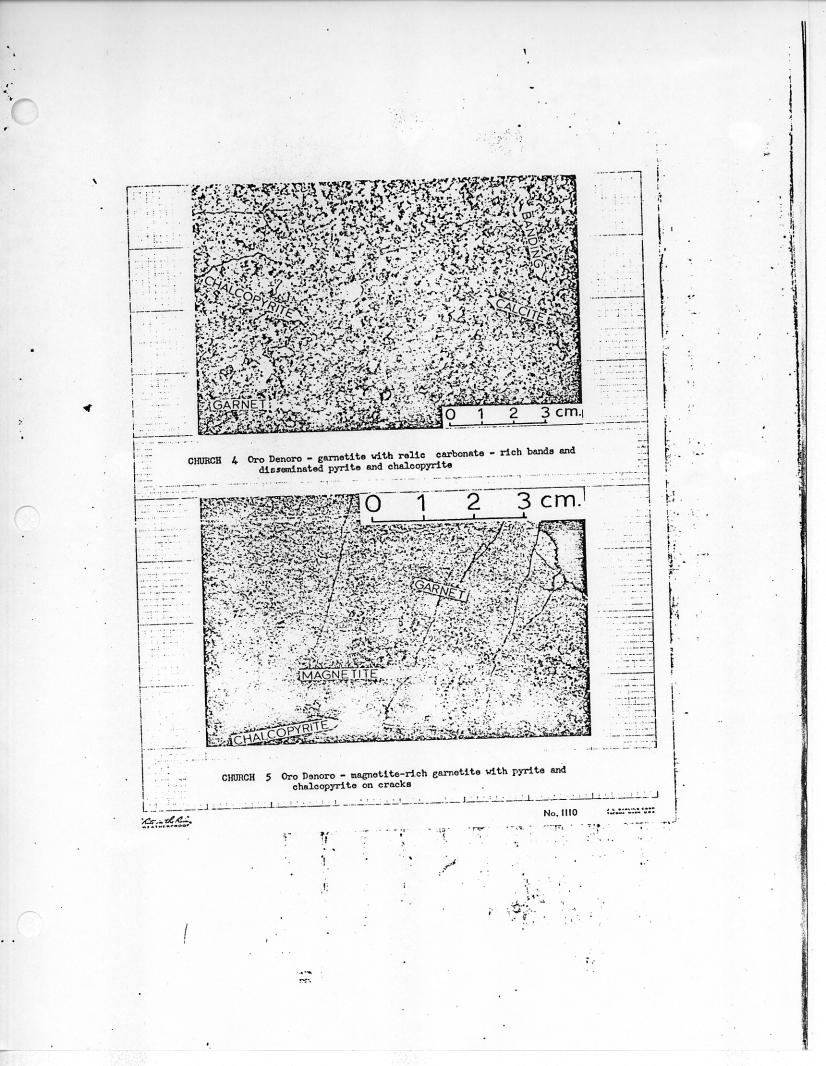
CHURCH	1	Oro Denoro - view of open pit area looking south
CHURCH	2	Oro Denoro - banded garnet magnetite, epidote skarn
CHURCH	3	Oro Denoro - garnetite with calcite and pyrite fillings
CHURCH	4	Oro Denoro - garnetite with relict carbonate - rich bands and disseminated pyrite and chalcopyrite
CHURCH	5	Oro Denoro - magnetite-rich garnetite with pyrite and chalcopyrite on cracks

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le la Key to Analyson - Amphibolite, Knob Hill Group, Deadman Hillaren 2 - Sharpstone conformerste, Brosklyn Formation, Desdman Hill 3 - Baralt lova, Shalt Formation, near the junction of the Phoenix road and Highway 3 just degoed the southeast comer of the mag-area. 4 - Skarn from No.1- Guarry of the Gro Denere marie 5 - Mafis phonolite, Marron Formation, Deadinon H.11 t- Lion Crek granocliorite, just south of the 7 - Emma porphyry, west of the One Denore pit 8 - Puloshite dyke, No. 1. Querry of the One Denne mine. 9 - Corgeil monzodiorite, northmest of the

		•	Tal	le.	1				1	
•	,	2	TABL	E OF C	HEMICAL	ANALYSE	rs 7	8	9	-
Oxides	Recalcul	lated to ,	100 -			,				
5:02	49.66	75.60	52.28	36.71	56.21	66.44	65.71	60.07	51.95	: .
Ti 02	1.05	0.52	0.96	0.11	0.97	0.42	.0.47	0.90	1.13	1.1
Al, 0;	17.21	9.98	17.83	5.69	17.41	16.22	16.12	15.84	14.84	
FezO;	3.11	1.33	9.31	23.09	4.85	2.00	1.02	2.58	2.31	
FLO	8.83	3.30	1.36	1.16	1.94	2.25	3.89	3.66	4.22	1.
MnO	0.22	0.12	0.15	0.39	0.08	0.06	0.14	0.12	0.11	11
MgO	7.07	2.94	8.17	0.64	2.24	2.00	2.52	5.12	4.53	
CaO	8.51	3.86	4.85	· 32.20	4.37	4.46	4.49	4.12	5.33	
Na,0	2.17	1.42_	4.80.	0.01	5.61	3.46	4.19	4.73	3.03	
K,0	2.17	0.93	0.29	• • -	4.32	2.69	1.44	2.67	4.55	
	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	
Oxides	as Dete	rmined						·		
+ +1/20	2.11	1.14	3.49	0.28	2.77	0.87	1.21	1.98	1.27	
- H/20	0.24	0.27	0.35	0.11	1.25	0.24	0.19	0.78	0.77	
CO2	0.60	0.15	0.89	1.67	1.35	0.45	1.19	0.91	1.21	
P205	0.18	0.32	0.39	0.48	0.71	0.18	0.30	0.28	0.41	
S	0.01	0.02	0.01	0.11	0.01	0.02	0.06	0.02	0.02	
.S-0 ;	0.03	0.02	0.09	0.006	0.30	0.08	0.09	0.13	0.12	
BaO	0.08	0.06	0.02	0.001	0.21	0.12	0.11	0.11	0.09 .	

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0120 DEVIDEO

CHURCH



