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Referee: *W. H. Miller*

DRE-FORMING PROCESSES <sup>IN</sup> (~~AT THE~~) HIGHLAND  
VALLEY PORPHYRY COPPER DEPOSITS, BRITISH  
COLUMBIA (~~A REVIEW~~)

BY

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HIGHLAND VALLEY  
ORE-FORMING PROCESSES  
PORPHYRY COPPER  
(and metal)

ABSTRACT

The Guichon Creek batholith located in south-central British Columbia is host to several producing and <sup>potentially</sup> ~~pre-~~producing porphyry copper deposits with an aggregate tonnage in excess of 2 billion tons (~~of material~~) grading approximately 0.4% Cu equivalent. (~~Most of~~ these deposits are <sup>within or close</sup> confined to the younger and more felsic units of the batholith in the Highland Valley district. Ore-forming processes are reviewed in light of (~~known and newly-acquired~~) geochemical and isotopic data.

not clear -  
specify nearly  
simultaneous  
emplacement and  
mineralization  
or ?

Published results of <sup>P</sup>radiometric and stable isotope studies indicate a close relationship between mineralization and emplacement of the Guichon Creek batholith. Relatively low K, Rb and Rb/Sr values, and high K/Rb ratios in batholithic rocks (<sup>as well as</sup> ~~are consistent with~~) primitive Sr and stable isotope data ~~and~~ suggest derivation of the Guichon Creek magma and associated metals from a subcrustal source.

Distribution of trace elements especially Cu in rocks <sup>of the batholith</sup> and minerals is however inconsistent with the prevalent hypothesis that ore metals were derived and concentrated by crystallization-fractionation (~~of a~~ <sup>from the</sup> Cu-rich) magma. On the contrary, it is argued that the Guichon Creek magma became impoverished in Cu as a result of fractionation. Thus, the role of the Guichon Creek magma in ore-forming processes is envisaged as one of structural control in channeling ore metals to the surface and providing structural openings for localization, <sup>of ore deposits</sup>

(A)

rather than a direct source of metals. The nature of epigenetic alteration-mineralization processes is also examined in light of bedrock and mineral geochemistry.

Available evidence suggests that altered wall rocks <sup>are</sup> ~~may~~ not ~~be~~ the source of ore metals.

## INTRODUCTION

In recent years, numerous genetic models have been proposed for porphyry copper deposits (Burnham, 1967; Meyer and Hemley, 1967; Fournier, 1967; Nielsen, 1968; Lowell and Guilbert, 1970; White, 1968; Philips, 1973). Most of these models <sup>were not supported by studies</sup> have not benefited from results of bedrock geochemistry, which in conjunction with experimental <sup>lead to a better</sup> studies are crucial to the understanding of chemical aspects of ore-forming processes in porphyry coppers. The purpose of this paper is to review ore-forming processes at Highland Valley, and on the basis of newly available geochemical data (Olade, 1974; Olade and Fletcher, 1975a, 1975b) <sup>to</sup> speculate further on the origin of the mineral deposits.

Most <sup>of</sup> the genetic models presented ~~for~~ porphyry copper deposits recognize the importance of magmatism in hydrothermal processes; the main differences <sup>between models</sup> are in the depth of intrusion, the timing of hydrothermal processes and source of mineralizing fluids (Lowell and Guilbert, 1970). In the <sup>orthomagmatic</sup> models (Burnham, 1967; Nielsen, 1968), an aqueous-rich volatile phase is released from the magma when internal vapour pressure exceeds lithostatic pressure, or when the intrusive system is subjected to external stresses. At the other end of the <sup>ore-genetic</sup> 'spectrum' <sup>(to the</sup> in orthomagmatic models, <sup>White, (1968)</sup> postulates an almost

hydrothermal solutions which may be of connate and/or meteoric hydrothermal solutions subject to convective movement as a result of processes by heat generated by subjacent intrusions. In this model, the pluton plays a passive role in mineralizing processes.

*is postulated. The source*  
*connate*  
*origin - arc*  
*of this type*  
*ation*

GEOLOGIC SETTING OF GUICHON CREEK BATHOLITH

*pic.*  
 The Geology of the Guichon Creek batholith has been described in detail by Northcote (1969) McMillan (1972) and Hylands (1972). *It is of late Triassic age and* The Triassic batholith intrudes sedimentary and volcanic rocks of the Permian Cache Creek Group and Upper Triassic Nicola Group, *It is unconformably overlain* and is overlain unconformably by Middle Jurassic to Tertiary sedimentary and volcanic rocks. The *pluton consists* pluton is composed of several concentric intrusive phases that range in composition from hybrid diorite *inside towards* at the outer margins to porphyritic quartz monzonite *in the center* at the core (Northcote, 1969). Most of the major porphyry copper deposits are associated with the younger *and* central phases of the batholith (Fig. 1).

CHARACTERISTICS OF MINERALIZATION IN THE BATHOLITH  
REVIEW OF EVIDENCE

Various lines of evidence suggest close relationships between mineralization at Highland Valley and evolution of the Guichon Creek batholith (Northcote, 1969; Grabc and White 1971).

Firstly, most of the major porphyry copper deposits are spatially associated with the younger and most felsic <sup>components</sup> ~~parts~~ of the batholith, the Bethsaida and Bethlehem Phases. Secondly, isotopic age determinations indicate <sup>a</sup> close temporal relationship between magmatism and hydrothermal processes. <sup>Results of K-Ar age determinations on <sup>fresh biotites (Northcott, 1971) and</sup> hydrothermal sericites and biotites (Blanchflower, 1972; Jones et al., 1972; Dirom, 1965) indicate that, within limits of analytical error, ~~mineralization and~~ <sup>and mineralization</sup> emplacement of the batholith were contemporaneous.</sup>

Origin of Guichon Creek Magma

only two references  
better to say  
"various"

Numerous workers have shown that K/Rb ratios set (important) constraints on (the) source materials of igneous masses (Hurley, 1968; Culbert, 1972). ~~Results of~~ <sup>Regional</sup> geochemistry (Olade, 1974) indicate that K/Rb ratios in rocks of the Guichon Creek batholith are relatively high (mean = 358) and largely outside the <sup>(reference)</sup> limit considered normal for continental plutonic rocks (Fig. 2). Furthermore, (the) Rb/Sr ratios plot in the region of basalts and andesites (Fig. 3) and the mean Rb/Sr ratio of 0.05 is one-fifth the value cited for sialic crust by Faure and Hurley (1963). Compared to other Mesozoic plutons in the Intermontane Structural Belt (Table 1), the Guichon Creek batholith is relatively impoverished in Rb and K, and characterized by higher K/Rb and lower Rb/Sr ratios. However, values obtained for the Guichon Creek batholith are similar

specify  
limit

Why not add these data to Table 1?

to those reported by Culbert (1972) for the Coast Mountains batholith of the Coast Mountains Structural Belt.

(The) <sup>the</sup> relatively high K/Rb and low Rb/Sr ratios in rocks of Guichon Creek batholith suggest derivation from a subcrustal source region <sup>which is</sup> depleted in alkalis and enriched in Sr, most <sup>likely from either</sup> ~~probably from~~ subducted oceanic crust or upper mantle. This interpretation is consistent with the primitive initial Sr isotopic ratio ( $Sr^{87}/Sr^{86} = 0.7037$ ) reported by Christmas et al., (1969). Furthermore, ~~results of~~ sulfur isotopes in hydrothermal sulfides and sulfates, <sup>as well as</sup> and deuterium and oxygen isotopes in hydrothermal sericites and kaolinities (Field et al., 1973; Jones et al., 1972; Sheppard et al., 1969) suggest a subcrustal source for mineralizing solutions and <sup>their</sup> associated metals.

Monger et al. (1972) and Dercourt (1972) ~~have~~ presented tectonic models for the evolution of the Canadian Cordillera. ~~which~~ <sup>They</sup> suggest that the Intermontane Belt, comprising extensive andesitic volcanic rocks and calc-alkaline plutons (including the Guichon Creek batholith), was the site of an ancient island arc <sup>which was</sup> generated by subduction of oceanic crust of the Pacific Plate beneath continental crust of the overriding North American Plate during ~~the~~ <sup>time.</sup> Mesozoic. In accordance with this model and ~~studies by Hetherington and Dickinson (1969) on~~ <sup>\* replace with comment in referee note</sup> ~~other island arc systems,~~ the relatively early Mesozoic age (200 m.y.) of the Guichon Creek batholith and its low K<sub>2</sub>O

content (mean = 1.85%) suggest derivation at relatively shallow <sup>(150 to 200 km) in</sup> depths ~~from~~ the subduction zone <sup>and 350 to 400 km)</sup> close to the Triassic 'trench'.

(C)

(C)

Emplacement of the Guichon Creek Batholith

Northcote (1969) has presented geologic evidence which suggests that ~~the~~ older intrusive units within the batholith were emplaced under mesozonal conditions, whereas ~~the~~ younger units that are spatially associated with mineralization were emplaced at shallower levels in the crust. Particularly towards the close of intrusive activity, volatile pressure ~~did~~ exceeded load pressure and tensile strength of the confining rocks, <sup>Consequently,</sup> ~~resulting in the development of~~ <sup>were developed</sup> breccia pipes as a result of subvolcanic explosions in an epizonal environment (White et al., 1957).

*Core, 1966  
Breccias, Porphyries & Cu Minerals in H.V.  
Western Montana*

Distribution of Trace Elements

*Bracc 43 to 65*

Within the batholith as a whole, Cu concentrations generally decrease from more than 300 ppm in the relatively older mafic units to less than 10 ppm in the younger felsic rocks at the core (Fig. 4). <sup>Similarly,</sup> ~~Similar trends are shown by~~ Zn, Mn, Ti, V, Ni and Co ~~which~~ are all relatively depleted in the youngest, most felsic phases of the batholith. However, unlike Cu, these elements are strongly correlated with Fe and Mg ( $r = 0.7$ ), <sup>Apparently they substitute for Fe</sup> ~~which they are assumed to substitute for in~~ the lattices of feric minerals (Ulade and Fletcher, 1975b).

*very different  
to Bracc's  
figure! yet the  
10 ppm figure  
agrees with  
him.*

The <sup>weak</sup> correlation of Cu with Fe and Mg ( $r = 0.5$ ) and its

*r = 0.5 for both?*



relative ease of extraction with a sulphide selective leach (Ojade and Fletcher, 1974) suggests that copper sulfides account for a significant proportion of the Cu in "unmineralized" samples.

Nature of Alteration - Mineralization Processes

Extensive wall-rock alteration, <sup>m</sup> that is so characteristic of porphyry copper deposits constitutes the most visible evidence of interaction between host rocks and hydrothermal solutions. Meyer and Hemley (1967) among others, ~~have~~ <sup>that</sup> demonstrated the close temporal and genetic relationships ~~exist~~ between sulphide deposition and wall-rock alteration at porphyry copper deposits.

Mineralogy of <sup>minerals in</sup> alteration assemblages at Highland Valley deposits provides evidence of <sup>about the</sup> the composition of mineralizing fluids. All the ~~deposits of the~~ <sup>deposits</sup> Highland Valley contain sericite alteration either in association with kaolinite, quartz or K-feldspar. Argillization and sericitization of wall rocks require <sup>conditions of</sup> slight to moderate acidity (pH < 6) whereas abundant K-feldspar <sup>requires less acid conditions (pH > 7)</sup> suggests pH exceeding 7 (Barnes and Czamanske, 1967). Cross-cutting vein relationships suggest that K-feldspar with or without quartz is generally early in the paragenetic sequence, and followed by sericite and argillic veins or selvages. This sequence suggests <sup>became more acid as the</sup> (increasing acidity of) hydrothermal fluids with increasing <sup>evolved.</sup> evolution. However, at Valley Copper, K-feldspar envelopes

Jones sug  
gets initial  
pH ~ 1.5  
dropping to  
2.5 to 3 during  
main stage  
mineralization  
then pH ~ 4  
when Kspar  
deposited  
at Valley  
Copper

occur around sericite veins and <sup>are</sup> in equilibrium with kaolinite.

This relationship, which is contrary to the stability-field relationships established for these minerals by Hemley and Jones (1964), is attributed to <sup>an influx of ore-forming fluids with</sup> a resurgence of <sup>abnormally</sup> high silica activities in <sup>ore-forming fluids</sup> ~~ore-forming fluids~~ <sup>shifting</sup> the mineral stability field to higher pH levels.

~~Results of~~ <sup>B</sup> Bedrock and mineral geochemistry (Olade, 1974) <sup>document that</sup> suggest that widespread chemical changes <sup>occurred</sup> in wall rocks <sup>are</sup> intimately associated with mineralization and hydrothermal alteration. Each deposit is characterized by central mineralized zones in which metasomatic activity <sup>was</sup> is most intense.

In zones of intense argillic and phyllic alteration at Valley Copper, Lornex and Highmont, the base elements Ca, Na, Sr, Ba, Zn, Mn, Mg and Fe are depleted, whereas in potassic zones at JA, Lornex and Valley Copper K, Rb and Ba are relatively enriched. Calculations of chemical gain and loss of principal rock constituents <sup>ferrous elements</sup> through alteration and mineralization at Valley Copper (Fig 5), suggest that in quartz-sericite and potassic zones, Ca, Mg, Fe, Na and Al are removed and <sup>whereas</sup> K, Si and S <sup>are</sup> added (for method of calculation, see Gresens, 1967). The <sup>(obvious)</sup> depletion of base cations in mineralized and altered zones is attributed to the breakdown of ferromagnesian minerals <sup>and</sup> plagioclase to sericite and kaolinite. Incipient stages of the above <sup>alteration</sup> process are demonstrated by results of mineral analysis<sup>s</sup>. Zn, Mn, and

*possibly Ni also?*

Co levels in biotites and magnetites are consistently lower in mineralized and altered than in fresh samples (Table 2). Cu and S concentrations, though erratic, are highest in zones of intense alteration and metallization, *and decrease* decreasing outwards to background levels in fresh unmineralized host rocks.

DISCUSSION

The following modes of origin have been proposed for porphyry copper deposits, hence are relevant to the *interpretation* genesis of the Highland Valley deposits.

(i) *Derivation* Extraction of ore metals by leaching of wall rocks by convecting meteoric waters ~~as proposed by~~ (White, 1968), or by deuteric alteration ~~as suggested by~~ (Putman, 1972).

(ii) Derivation of ore metals by assimilation of metal-rich country rocks (Schau, 1970).

(iii) *Derivation of ore-forming fluids from* Derivation and concentration of aqueous-rich, *magmas by a process of fractionation during* metal-bearing fluids from magmas during their emplacement *crystallization* and crystallization-fractionation (Nielsen, 1968; Brabec and White, 1971; Graybeal, 1973).

(iv) Generation of ore-forming fluids from similar source regions as the magmas; the role of magmas being one of structural control rather than *acting as source* ~~source~~ metals (Noble, 1970, Sillitoe, 1972; Mitchell and Garson, 1972; Wright and McCurry, 1973).

On the basis of variations of Cu contents in the Guichon Creek batholith, Brabec and White (1971) postulated that the Highland Valley deposits were derived and concentrated during emplacement and crystallization of a <sup>1/2</sup> probably-Cu-rich Guichon Creek magma. In contrast Schau (1970) suggested that Cu in the batholith was generated <sup>by</sup> from the assimilation of a (supposed) Cu-rich Nicola volcanic country rocks. On the basis of Sr and other stable isotopes studies, Christmas et al. (1969) (have) proposed an upper mantle source for the Guichon Creek batholith and associated mineral deposits. These alternatives are now considered in light of <sup>regional</sup> regional and detailed bedrock, and mineral geochemistry <sup>analyses</sup>.

*I thought they only worked in Craigmont?*

The first hypothesis concerning derivation of metals from wall rocks is considered least likely because <sup>at the level of sampling,</sup> results of both regional and detailed bedrock geochemistry (Olade, 1974; Olade and Fletcher, 1975b) <sup>adjacent to</sup> around mineralized zones <sup>showed</sup> indicate that no zone of Cu and/or S depletion surrounds the orebodies. <sup>Cu and S were</sup> (at the level of sampling) The possibility that these elements could be extracted from <sup>derived from wall rocks</sup> channelways <sup>at</sup> at greater depths is not ruled out. Moreover, results of mineral geochemistry provides no evidence of obvious leaching of Cu from biotites.

Schau (1970) suggested that ore metals were derived by assimilation of Nicola volcanic rocks. Subsequently Brabec and White (1971) criticized this hypothesis by demonstrating that the <sup>border</sup> Hybrid Phase, the most contaminated unit within the batholith, is not significantly higher in Cu than uncontaminated <sup>has approximately the same Cu content as</sup>

*↑ less than Guichon, more than Chataway*

*Varieties*

(D)

rocks of the Guichon and Chataway Phases. Brabec (1970) further suggests that the relatively high Cu levels <sup>of</sup> in the batholith would require selective assimilation of this metal <sup>of a very</sup> from a large volume of country rocks. Field evidence <sup>does</sup> not support a large-scale contamination of the batholith <sup>Only its</sup> beyond the outer margins <sup>show evidence</sup> of country rock assimilation (Northcote, 1969).

Available geochemical data are not consistent with the hypothesis of Brabec and White (1971), <sup>because:</sup> since

(i) Cu, together with Zn, Mn, Ti, V, Ni, Co, Fe and Mg generally decreases with increasing <sup>differentiation</sup> fractionation or felsic composition of intrusive units. This geochemical pattern simply reflects normal differentiation trends observed in unmineralized intrusions. Sheraton and Black (1973), <sup>who</sup> investigating trace element geochemistry of <sup>granitic intrusions unmineralized with Cu deposits?</sup> granitic intrusions unmineralized with respect to Cu, found that Cu concentrations decreased from more than 40 ppm in granodiorite to less than 5 ppm in more differentiated granites. In contrast, studies on intrusions that are known to have generated immiscible sulphide phases such as the Skaergaard <sup>intrusion</sup> (Wager and Brown, 1967), and mineralized Laramide intrusions in Arizona (Graybeal, 1973), Cu contents of bedrock and mineral constituents generally increase with fractionation until Cu separates from the melt as an immiscible sulphide phase. <sup>then decreases sharply.</sup> Graybeal

(E)

Confusing, was he measuring real rocks and comparing Cu in bi + bb to

(1973), investigating the partitioning of Cu between co-existing biotite and hornblende found that, under equilibrium conditions, higher concentration of Cu within the magma was reflected by higher <sup>Cu</sup> concentrations in the <sup>two</sup> mineral phases. In the Guichon

Whole rock copper? or was he making the magma? is doing "bomb" work?

Creek batholith, results of Cu determinations in <sup>18</sup>biotites and <sup>4</sup>hornblendes <sup>concentrates</sup> (Brabec, 1970) suggest no appreciable variations throughout the batholith. From the foregoing discussions it is <sup>concluded</sup> apparent that geochemical data <sup>are</sup> do not <sup>consistent with</sup> support the hypothesis that ore metals <sup>in</sup> at Highland Valley <sup>deposits</sup> were derived by differentiation of a Cu-rich Guichon Creek magma. On the contrary, it is argued that the Guichon Creek magma became increasingly impoverished in Cu as a result of differentiation.

An hypothesis, which regards mineralization as an independent by-product of magma generation rather than a direct result of differentiation processes, is consistent with geochemical data and contemporary ideas <sup>about</sup> of plate tectonics and ore genesis.

Nevertheless, it must be emphasized that differentiation processes within a magma <sup>provides the right</sup> provide the right chemical <sup>formation</sup> and <sup>setting</sup> emplacement of the magma <sup>deposits.</sup> physical environment for localization of ore metals.

High <sup>ratios</sup> K/Rb and Sr values, and low <sup>and K values</sup> Rb, <sup>and Rb/Sr ratios as well as low Sr</sup> <sup>87</sup>/<sub>Sr</sub> <sup>86</sup>/<sub>Sr</sub> and Sr

isotopic ratios are consistent with derivation of Guichon Creek magma from a deep-seated source, <sup>either</sup> (most) probably <sup>(Jones et al, 1972)</sup> subducted oceanic crust or upper mantle. \* Results of sulphur, oxygen and deuterium isotopes suggest a similar deep-seated source for mineralizing solutions and ore metals.

Because of the temporal and spatial relationships between mineralization and magmatism, it is logical to presume that ore metals at the Highland Valley deposits were derived from a metal-rich portion of the subducted oceanic crust from which the Guichon Creek magma was generated.

proof that the magma was not generated by partial melting of upper mantle?

(F)

Sillitoe (1972) ~~(has)~~ demonstrated that there is enough Cu in oceanic basalts to generate metals in ore deposits. The ore metals derived from partial melting of subducted oceanic crust probably occur in a <sup>which is</sup> phase independent of the magma. Thus the ~~role of the magma is believed to be one of~~ <sup>magma has a structural role and opens access</sup> structural control in channeling ore metals to <sup>more up to</sup> crustal levels (Noble, 1970). <sup>Emplacement and</sup> Nevertheless, differentiation of the magma would provided volatiles and structural openings, such as fractures, <sup>and</sup> breccia zones, <sup>to</sup> dyke swarms that facilitated ~~(the)~~ extraction of metals from the system and concentration as ore deposits.

*not necessarily permeable zones although the fractures they occupy were.*

Fig. 6 shows a ~~(comprehensive)~~ model for ~~(the)~~ evolution of ~~(the)~~ ore-forming fluids at the hydrothermal stage <sup>in Highland Valley.</sup> The close spatial relationship between porphyry dykes or dyke swarms and ore deposits <sup>at</sup> Highland Valley, suggests that the fracture systems <sup>which controlled dyke emplacement also</sup> porphyries served as high-level structural 'outlets' for mineralizing solutions. The presence of saline fluid inclusions in quartz veins at Valley Copper Lornex and Highmont (R.D. Barton, pers. comm.) <sup>as well as</sup> and enhanced values of S, F, Cl and S in ore zones suggest that the mineralizing fluids contained HCl, H<sub>3</sub>BO<sub>3</sub>, HF, H<sub>2</sub>S, H<sub>2</sub>SO<sub>4</sub> and other volatile elements. Late stage differentiation products, such as K<sup>+</sup>, SiO<sub>2</sub>, Rb<sup>+</sup> and Na<sup>+</sup> were probably present. Extensive argillic and sericite alteration found around the deposits require that ore solutions be slightly to moderately acidic, and contain abundant H<sup>+</sup>, probably derived from dissociated H<sub>2</sub>O and H<sub>2</sub>S ~~(present)~~ in ~~the~~ the

④

juvenile fluids, or ~~(by admixture with convecting)~~ meteoric waters, ~~(generated by heat from the porphyry dykes or stocks),~~

Helgeson (1970) ~~(has)~~ presented thermodynamic data which <sup>suggest</sup> ~~(demonstrate that (all))~~ equilibria in hydrothermal systems can be represented in terms of the ratio of activities of cations in the aqueous phase to that of the hydrogen ion. Changes in base cation/H<sup>+</sup> activities <sup>in</sup> ~~as~~ ore-forming fluids <sup>as they pass through various</sup> transgress the alteration zones are portrayed in Fig. 6. The evolutionary

Which is which? →

<sup>paths followed by fluids with</sup> paths, designated 1, 2 and 3 in the diagram, represent <sup>states</sup> different degrees of equilibration between <sup>with the</sup> ore fluids and wall rock.

Formation of an early potassic zone <sup>containing</sup> (K-feldspar ± quartz ± sericite), that is commonly centred on porphyry dykes, requires a high base cation (K<sup>+</sup>, Na<sup>+</sup>)/H<sup>+</sup> activity ratio, <sup>This</sup> which could <sup>reflect</sup> result <sup>the</sup> from initial compositions of mineralizing fluids ~~(inherited from the magma)~~ or less <sup>likely, could</sup> probably be derived at depths by H<sup>+</sup> consuming and base cation-releasing equilibrium reactions, <sup>at depth</sup>

As ~~(the)~~ ore fluids rise and spread outwards they undergo adiabatic expansion, and <sup>as a result of</sup> ~~(in conjunction with)~~ reaction with wall rocks and/or mixing with meteoric waters, <sup>they</sup> cool, causing dissociation of the most acidic components. This dissociation provides most of the abundant H<sup>+</sup> required for hydrolytic base leaching within ~~(the)~~ quartz-sericite <sup>ic</sup> and argillic <sup>alteration</sup> zones, under acidic conditions. <sup>are</sup> The base cations (Mg<sup>++</sup>, Ca<sup>++</sup>, Fe<sup>++</sup>, Na<sup>+</sup>, Sr<sup>++</sup>, Ba<sup>+</sup>, Zn<sup>++</sup>, Mn<sup>++</sup>) released by leaching, <sup>enter</sup> ~~are taken into~~ the fluid and <sup>are</sup> transferred to the outlying metasomatic front (Korzhinskii, 1968), as the solutions are cooled and neutralized,



Changes in base cation/ $H^+$  activity ratios are generally accompanied by changes in pH and sulfur fugacity (Meyer and Hemley, 1967) which ultimately control sulfide deposition and zoning patterns. ~~This accounts for the~~ close association between sericite and argillic alteration which require  $H^+$  consumption in their formation and sulphide mineralization, <sup>results,</sup> as <sup>is</sup> amply demonstrated at Valley Copper, Lornex, Highmont and in parts of Bethlehem-JA.

From the foregoing discussion, it is apparent that regional <sup>and</sup> detailed bedrock and mineral geochemistry, ~~and~~ isotopic and tectonic evidence are consistent with the mode of origin proposed for the Guichon Creek batholith. Assuming the genetic model <sup>is</sup> correct, it has ~~far-reaching~~ implications in reconnaissance exploration for Cu deposits in calc-alkaline intrusions of

the Intermontane Belt. First, the apparent negative correlation between Cu contents and ore potential of the Guichon Creek batholith suggests that ore-bearing intrusions need not be enriched in Cu. Thus the suggestion by Warren and Delavault, (1960) that high Cu contents of intrusions reflect ore potential might not be generally applicable.

Secondly, if ore metals in the Guichon Creek batholith were derived from <sup>copper-rich</sup> subducted oceanic crust as an independent by-

product of magma generation, it is most <sup>probable</sup> plausible that other calc-alkaline plutonic and volcanic rocks <sup>in the Nicola belt which are</sup> of similar age as the Guichon Creek batholith might originate from <sup>similar</sup> the same

The batholith as a whole is copper positive - only the phases associated with the deposits are copper deficient

H

I

metal-rich (~~portion of~~) subducted oceanic crust. Such calc-alkaline intrusive and extrusive rocks within the Intermontane Belt <sup>could</sup> be identified by: (1) <sup>late Triassic to early Jurassic ages</sup> their ~~ages~~ (~~Late Triassic - Early Jurassic~~); (2) their low Rb, Rb/Sr and high K/Rb ratios; and (3) their <sup>low</sup> ~~high~~  $K_2O$  content which should reflect the relatively shallow depth of magma generation. Using the Guichon Creek batholith as a 'reference index', calc-alkaline intrusive and extrusive rocks which meet the above criteria might have (~~considerable~~) potential for further discoveries of porphyry Cu and/or massive sulfide deposits.

### CONCLUSIONS

Regional, <sup>and</sup> detailed bedrock and mineral geochemistry of the Guichon Creek batholith and associated mineralization is consistent with the hypothesis that ore metals did not arise as a direct result of differentiation processes within a Cu-rich magma, but rather as an independent by-product of magma generation from subducted oceanic crust, of probably amphibolite composition. Nevertheless, <sup>emplacement as well as</sup> chemical and mineral fractionation within the Guichon Creek magma led to the (development of) increased volatile contents and pressures that provided suitable chemical and structural environments

CMIT  
never  
previously  
mentioned

for localization of ore deposits. Consequently not all  
*omit* ore-bearing plutons need be enriched in Cu.

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? not seen in the text

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Sillitoe 1972

White Thompson  
& McT 1957

WJM Comments

- (A) Abstract bottom 4 lines ..... I think you are suggesting that the mineralizing solutions were either (1) emplaced with the magma as an immiscible phase or (2) that they followed behind the magma but used the same channelways. It is not clear to me which possibility you mean. If it is (1) it is difficult to envisage a magma structurally controlling channelways. It should also be made clear that the structural openings for localizing ores were in the crystallized carapace of the magma.
- (B) It would be better to be more specific.
- (C) You might be able to improve the depths and distances from trenches which I got from Hatherton and Dickson's curves by calculating K55 and K60 for Guichon data. There should also be a qualifying statement to the effect: "assuming conditions at the time the Guichon was emplaced were similar to
- (D) I strongly criticize your decision to change what has become standard usage (Northcote, 1969) regarding the Highland Valley PHASE and the Guichon and Chataway VARIETIES. I suggest you stick to standard usage even though it would necessitate changing the Legend of Figure 1.
- (E) Graybeal concluded that copper content was low in productive plutons because copper separated out into the volatile phase." At Skaergaard, copper in the rock and minerals was low after the sulphide phase separated. In the Guichon, Brabec shows copper increasing from Hybrid (57) to Guichon (65) then dropping (Chataway 43, Bethlehem 32 and Bethsaida 10). All figures are geometric means.
- (F) This statement must be preceded by some logical reason for rejecting derivation from a metal-rich upper mantle source. I would add a sentence in at the "\*" (p.11) saying:-

"In the Coast Mountains batholith Culbert (1972) has argued that similar Rb/K ratios indicate probable derivation of alkalis in the batholith at least in part from destruction of oceanic crust."



- (G) Your chemical gain calculations (p.7) are good evidence that K, Si, and S were added and enrichment shows Rb and Ba addition BUT Na is supposedly depleted (p.7).
- (H) At Valley in particular, argillic alteration is not closely associated with ore grade material. Argillic alteration at Valley is taken to be pervasive sericite and kaolinite development. Reference "Osatenko, M.J. and Jones, M.B., 1975 (?), Valley Copper Deposit : C.I.M. Special Volume 15, in preparation.
- (I) Does it really matter whether the sulphide phase was independent? So long as the magma and sulphides have the same general source area the rest of your argument follows logically.

FIGURE 1

- (a) HIGHMONT not HIGHMOUNT.
- (b) Add a 6 to the area of Bethsaida south of Highmont and east of Lornex Fault.
- (c) Are the two unnumbered areas near Ashcroft Barnes Lake and Willard Lake - if yes, remove them.
- (d) JA deposit not shown (label it J.A. when you add it so people unfamiliar with the region will know which of the Bethlehem deposits it is).
- (e) Suggest you follow standard usage and use:-
  - Highland Valley phase
  - Guichon Variety
  - Chataway Variety

FIGURE 2

- (a) Label the limits of normal continental plutonic rocks on the diagram itself.
- (b) Show the average, specify the standard deviation for Guichon rocks.
- (c) What is the significance of the  $r = 0.78$ . It is not mentioned in the text.

FIGURE 3

- (a) Significance of  $r = -0.52$ ?
- (b) If  $Rb/Sr = 0.01$  is typical for sialic crust, label it to that effect on the figure.
- (c) Dacite not Docite

FIGURE 6

The "model" is a representation of Highmont and possibly J.A. It is not representative of Bethlehem, Lornex or Valley Copper. I suggest you entitle the Figure:

Schematic model for chemical and mineral zoning of the Highmont deposits, Highland Valley area. Possible evolution of ore forming fluids is shown schematically."

NOTE: It would be better if the pH, ore fluid diagram, were separated from the other. I expected the "path" to reflect the underlying data but in fact it seems to be showing a change from the beginning to the end of the mineralization process.

TABLE 2

Terms "Fresh", "weakly mineralized" and "strongly mineralized" are confusing. Are the samples from porphyry deposits ..... if so specify this in the title, and perhaps use terms:-

Unmineralized rock  
Weakly mineralized zone  
Strongly mineralized zone

CAPTIONS TO TABLES

TABLE 1: Means and ranges of Rb, Sr, Rb/Sr, K/Rb and  $Sr^{87}/Sr^{86}$  ratios in some Mesozoic intrusions of the Cordilleran Intermontane Belt (After Peto, 1974).

TABLE 2: Means and ranges of some \*trace elements in biotites and magnetites from the Highland Valley (Values in ppm).

Archean?

TABLE 1: Means and ranges of Rb, Sr, Rb/Sr, K/Rb and Sr<sup>87</sup>/Sr<sup>86</sup> ratios in some Mesozoic intrusions of the Cordilleran Intermontane Belt (After Peto, 1974)

Intrusions	Rb (ppm)	Sr (ppm)	Rb/Sr	K/Rb	Sr <sup>87</sup> /Sr <sup>86</sup>	Age (m.y)
Guichon Creek batholith	35 ( 3-132)	686 (249-1000)	0.05 (0.004-0.321)	358 (132-1030)	0.7037	200 ± 5
Similkameen batholith	95 ( 52-152)	390 (147-639)	0.151 (0.081-1.01)	250 (172-309)	0.7060	183
Nelson batholith	- -	- -	0.175 (0.056-0.483)	- -	0.7069	171
Hogem batholith	80 ( 55-118)	730 (468-1520)	0.100 (0.041-0.125)	430 (322-502)	-	170
White Creek batholith	265 (196-357)	804 (435-1118)	0.412 (0.108-1.655)	- -	0.7250	126
Vernon batholith	- -	- -	1.42 (0.108-2.84)	- -	0.7064	55

TABLE 2: Means and ranges of some \* trace elements in biotites  
and magnetites from the Highland Valley (Values in ppm)

	Cu	Zn	Mn	Ni	Co
<u>BIOTITES</u>					
Fresh (10)	98 (39-163)	349 (172-931)	4195 (1515-11225)	40 (27-75)	56 (35-72)
Weakly Mineralized (8)	863 (474-2644)	292 (216-365)	3250 (2164-4388)	37 (25-54)	58 (47-76)
Strongly Mineralized (9)	2549 (531-5617)	275 (181-421)	3208 (1953-5821)	24 (12-44)	49 (42-54)
<u>MAGNETITES</u>					
Fresh (10)	67 (35-153)	68 (30-122)	-	-	40 (20-49)
Weakly Mineralized (8)	251 (93-846)	54 (38-64)	-	-	40 (34-59)
Strongly Mineralized (9)	576 (122-5451)	59 (39-73)	-	-	33 (24-37)

\* Atomic absorption analysis.

CAPTIONS TO FIGURES

- ✓ FIGURE 1: Location and general geology of Guichon Creek batholith (modified after McMillan, 1972).
- ✓ FIGURE 2: Plot of K versus Rb and K/Rb ratios in rocks of Guichon Creek batholith (Normal crustal ratios = 150-300).
- FIGURE 3: Relationship between Rb and Sr in rocks of Guichon Creek batholith (generalized geochemical relationships of Rb and Sr in certain rock types are shown ~~af~~ for comparison; after Hedge, 1966).

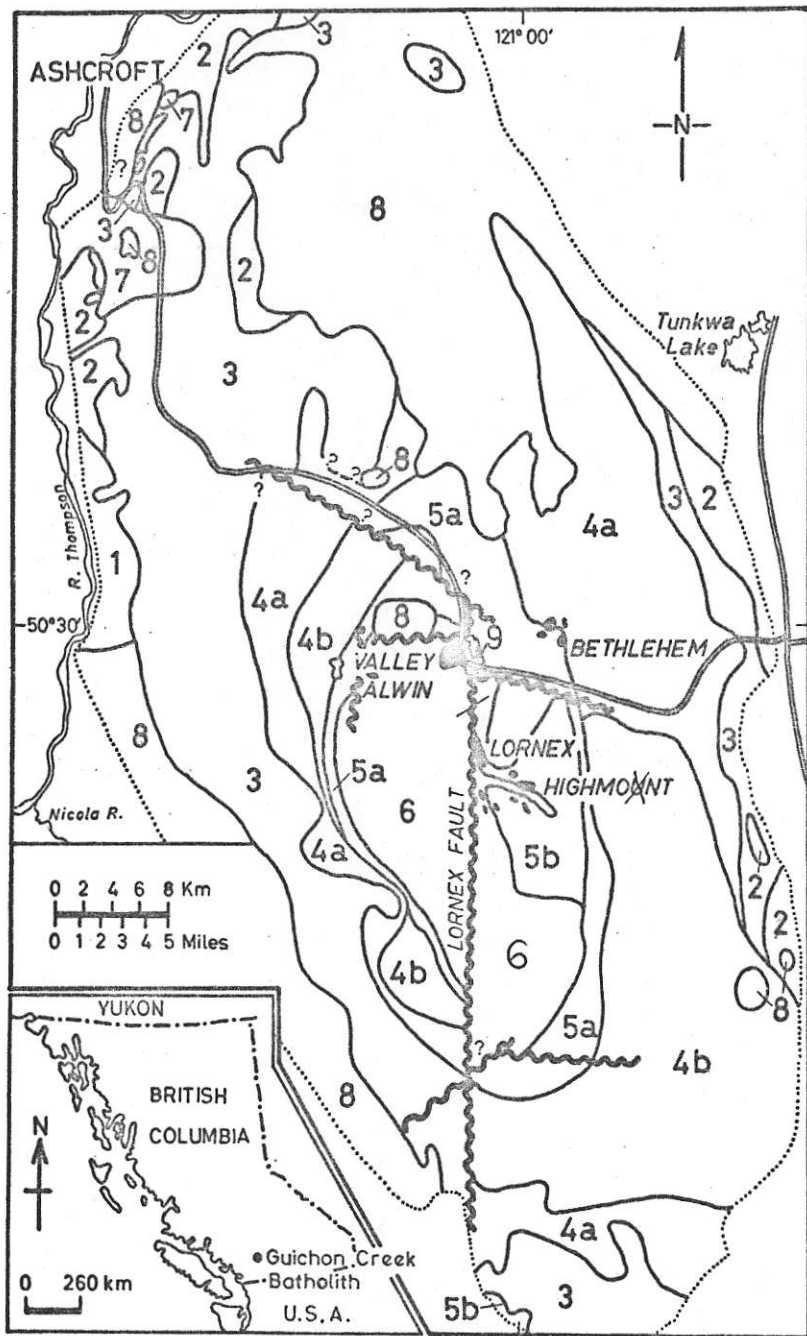
*add number of sample and grade as of list (or) to figure*

FIGURE 4: Distribution of Cu in Guichon Creek batholith in relation to Larsen differentiation index.

FIGURE 5: Gain and loss of principal rock constituents through <sup>various</sup> alteration and <sup>zones</sup> mineralization at Valley Copper (3600 level).

FIGURE 6: Schematic model for chemical and mineral zoning and evolution of ore-forming fluids <sup>in</sup> at Highland Valley porphyry copper ~~deposits~~.

*see sheet attached to Figure*



- LEGEND**
- LOWER OLIGOCENE
  - 9 Sedimentary & Volcanic Rocks
  - TERTIARY
  - 8 Volcanic Rocks
  - MIDDLE & UPPER JURASSIC
  - 7 Mainly Sedimentary Rocks
  - INTRUSIVE ROCKS
  - CORE UNIT
  - 6 Bethsaida & Gnawed Mt. Phases
  - INTERMEDIATE UNIT
  - 5b Skeena Phase
  - 5a Bethlehem Phase
  - HIGHLAND VALLEY UNIT
  - 4b Chataway Phase
  - 4a Guichon Phase
  - BORDER UNIT
  - 3 Hybrid Phase
  - PRE-INTRUSIVE ROCKS
  - UPPER TRIASSIC
  - 2 Nicola Volcanic & Sedimentary Rocks
  - PERMIAN
  - 1 Cache Creek Volcanic & Sedimentary Rocks
  - Ore Bodies
  - Geological Contact
  - Roads

Fig. 1: Location and General Geology of Guichon Creek Batholith. (Modified after McMillan, 1972)

Fig. 1



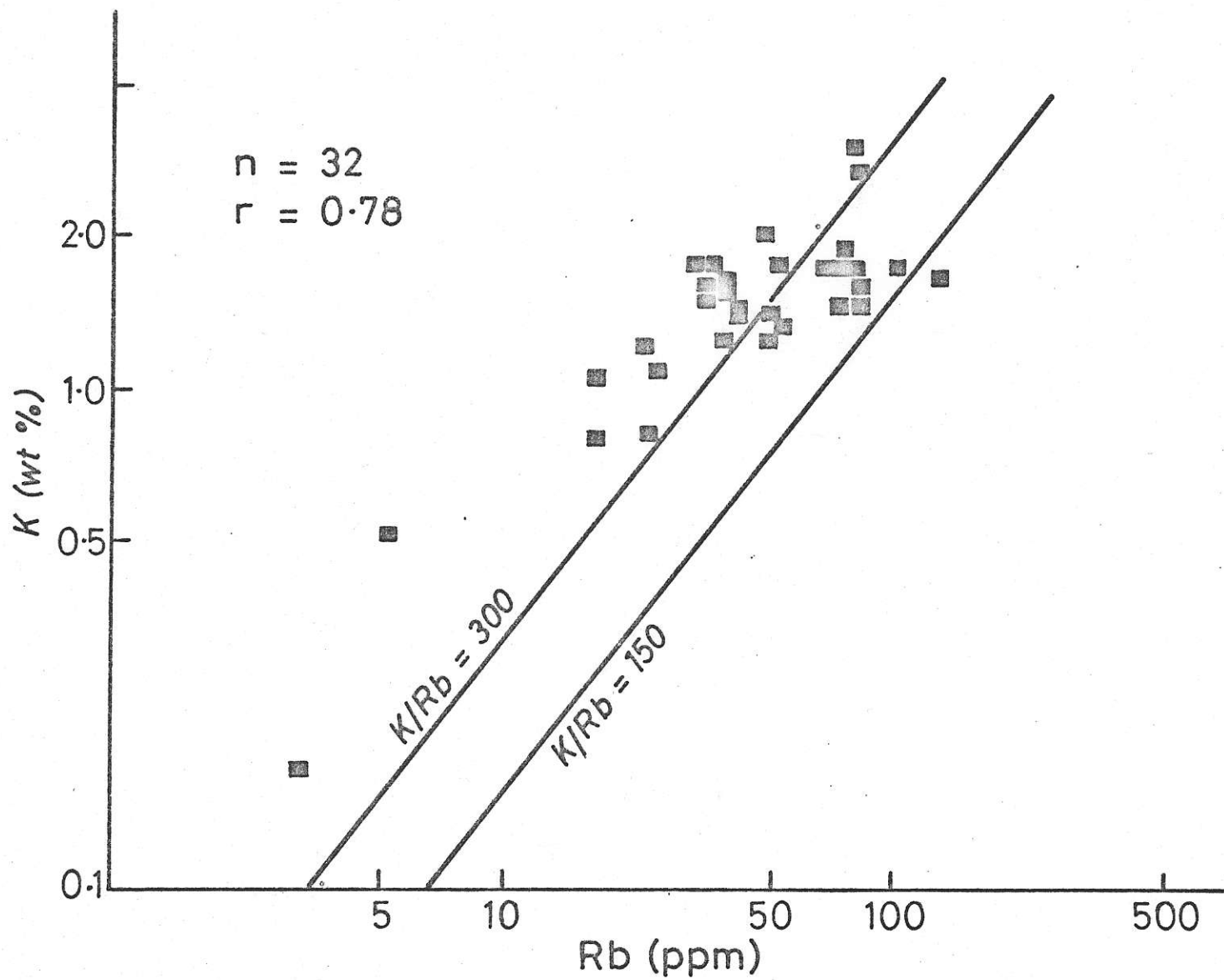


Fig. 2

• Guichon rocks

★ Average Sierra Nevada  
granitic rock

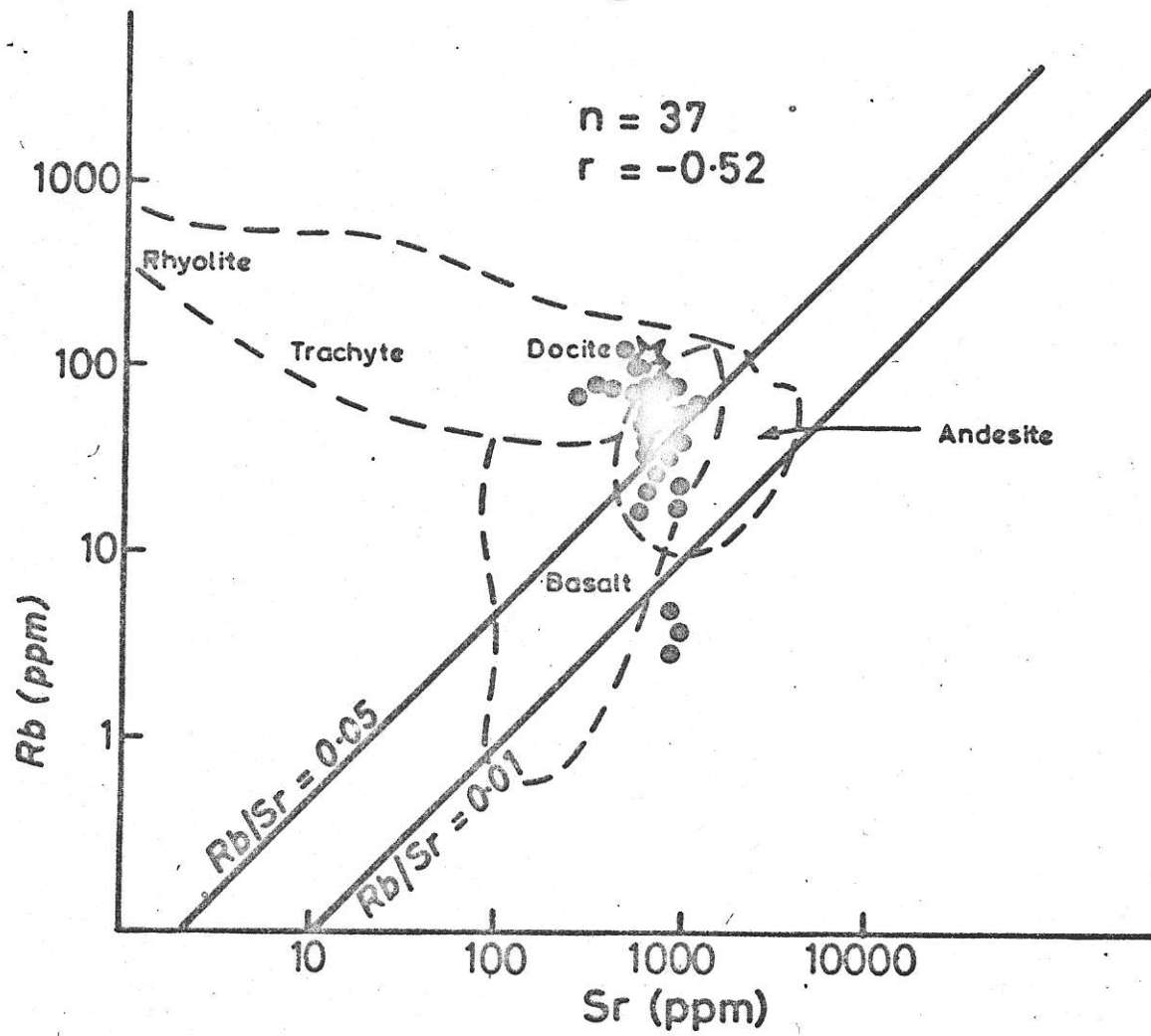


Fig. 3

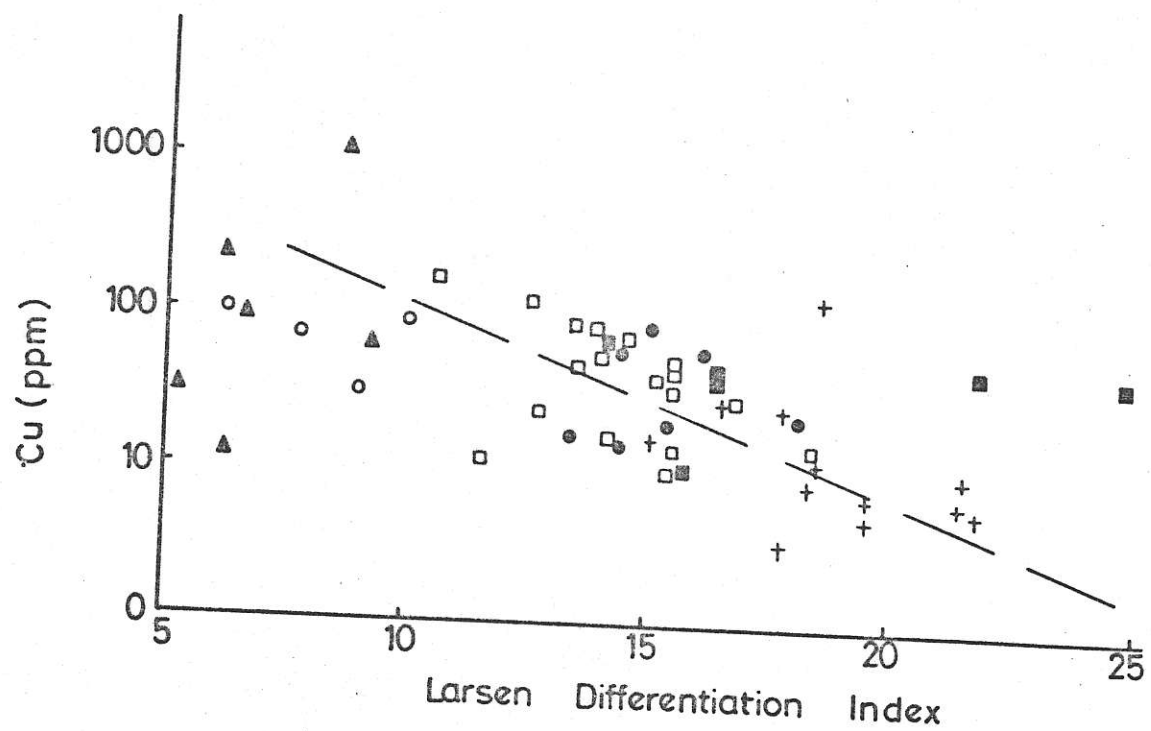


Fig 4

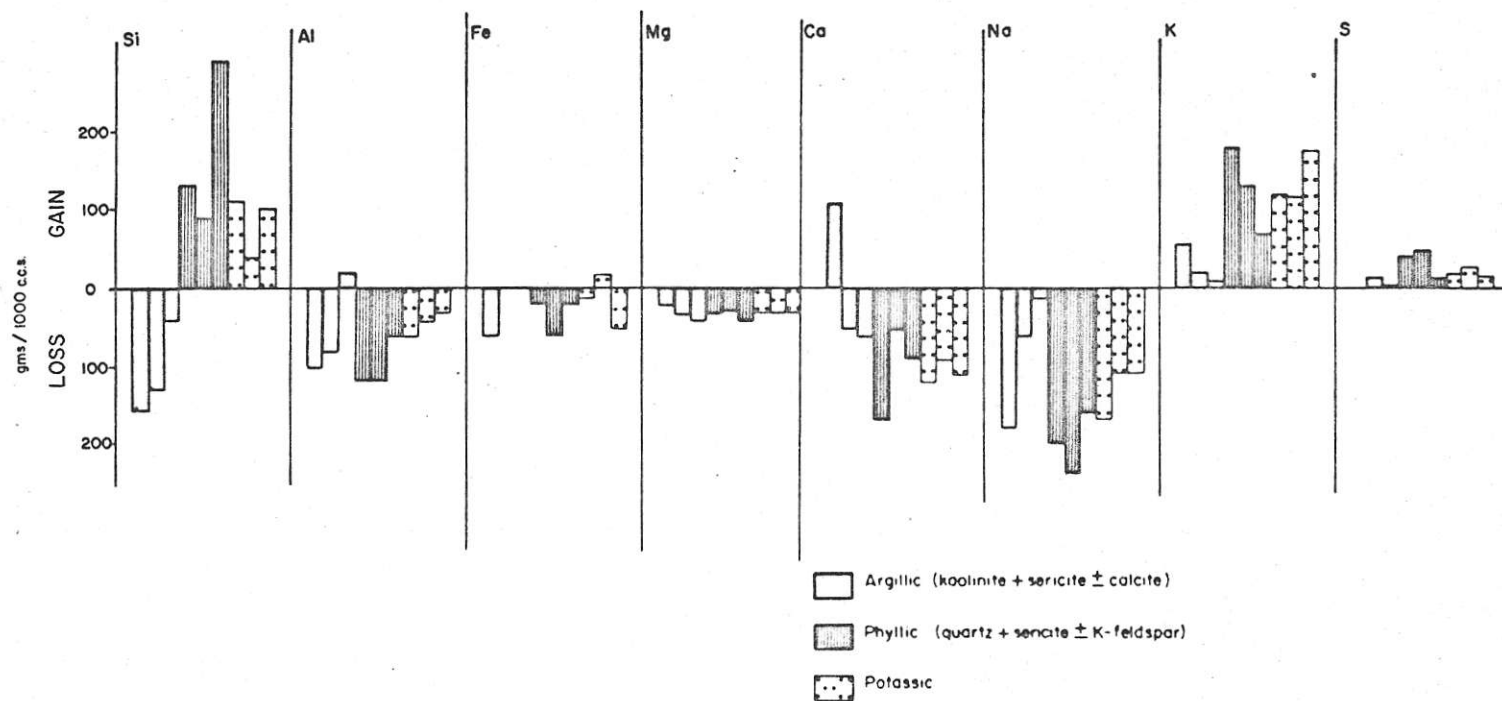
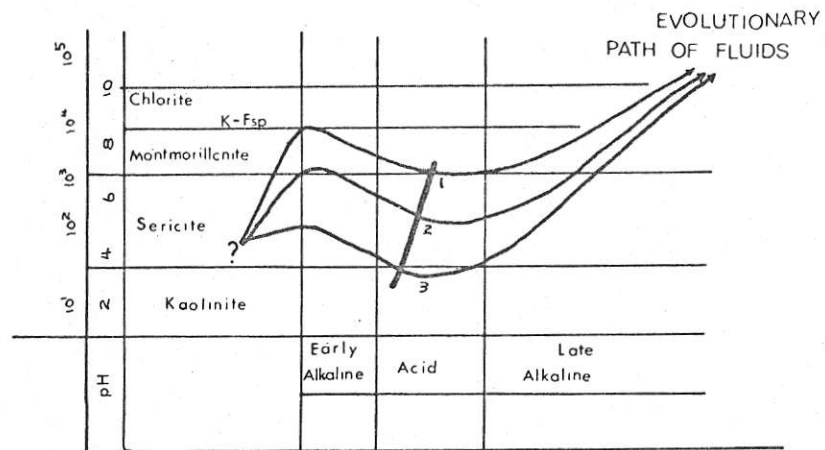
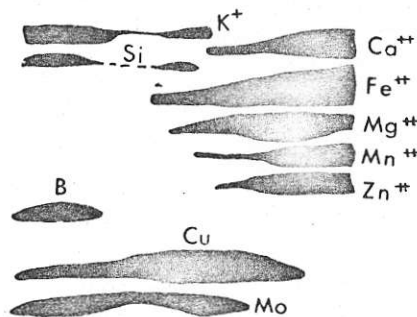


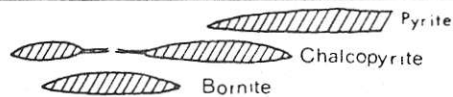
Fig 5



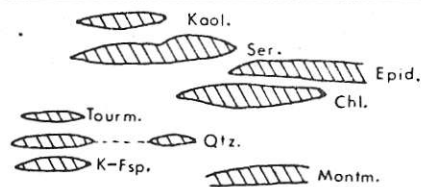
CHEMICAL  
ZONING



SULPHIDE  
ZONING



ALTERATION  
ZONING



ALTERATION TYPE	Pot.	Arg.	Qtz Ser	Propylitic
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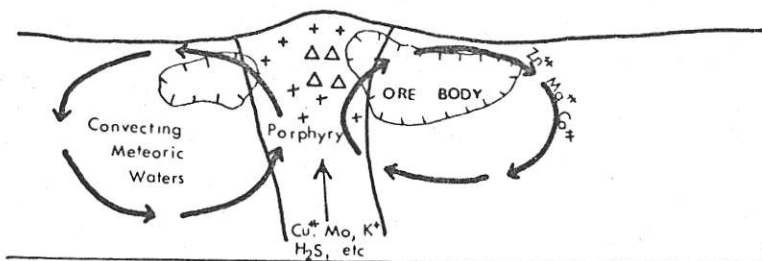


Fig 6