

Mineralographic Study of the  
"Maid of Erin" Copper Property  
Ore - Rainy Hollow, B.C.

Kenneth M. Dawson      Nov. 20/63  
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Mineralographic Study of the  
"Maid of Erin" Copper Property  
Haines - Alaska Cutoff  
Rainy Hollow B.C.

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UBC Mineralogy Report

Megascopic  
Description of Hand Specimens

There were 24 hand specimens studied in the tray. These were classified into three groups as follows:

- ① Predominantly metallic minerals - 13 samples
- ② Non-metallic minerals - 3 samples
- ③ Mixture of metallic and non metallic - 8 samples

1. Metallic Minerals

All samples showed a predominance of massive barnite with saffelite alteration. One sample showed several patches of chalcocite intimately associated with the barnite.

Another sample showed several grains of sphalerite about  $\frac{1}{4}$ " diameter with good cleavage. It was of the darker, non-red type.

All the metallic samples had small (1-3 mm) euhedral yellow-green garnets disseminated fairly evenly throughout the metallic matrix.

A little calcite was seen in the form of small rhombic cleavage fragments surrounding the garnet.

A small amount of quartz was seen in disseminated euhedral grains of 2-5 mm diam.

Alterations in the form of saffelite, malachite and a little azurite was noted on all specimens. Four samples had a soft, earthy limonitic alteration of yellow brown to grey-brown colour, possibly a copper salt.

average composition of 13 samples

Barnite 50%  
Garnet 25%  
Chalcocite 15%  
Sphalerite 8%  
Calcite 1%  
Quartz 1%

2. Now-metallic Hand Samples

The three samples in this group were almost uniform in composition. The predominant mineral was massive, deep-coloured garnet with fine twinning lamellae. Slender radiating prisms of pinkish brown tourmaline cut through the garnet. Some small anhedral grains of quartz were disseminated in the garnet in small amount. A few small grains of calcite surrounding smaller grains of chalcopyrite were seen in two of the specimens. One sample showed a little malachite stain around a zone of tourmaline prisms.

average composition of 3 samples:

Garnet	80%
Tourmaline	15%
Quartz	5%
Calcite	< 1%
Chalcopyrite	< 1%

3. Mixture of Metallic & Now-Metallic Minerals

The hand specimens all showed a coarse-grained, well-crystallized texture, with a larger variety of now-metallic gangue minerals than seen in other samples. One specimen was a homogeneous equigranular rock composed of calcite, barite and sphalerite. The calcite had been replaced by barite and sphalerite, then recrystallized into anhedral xls of 2-5 mm. The barite was disseminated in fine grained blades of 2-5 mm size some of which were connected. A little dark coloured sphalerite was seen also. Malachite and azurite stain was present.

The other six specimens were fairly uniform in texture and composition. Massive white quartz with massive light green vesuvianite intermixed formed the groundmass. Zoned, and randomly-distributed large tetragonal prisms ( $\frac{1}{2}$  inch) of green vesuvianite with massive bornite veinlets and masses around them were the most striking feature of these samples. Radiating prisms of pink tourmaline cut through the groundmass.

Specimen K442 had a matrix of light green to white vesuvianite in circular areas of slender radiating prisms. Bornite veinlets cut through this matrix. The thin section showed the fibrous aggregate nature of the groundmass of vesuvianite. Optical properties noted were parallel extinction, high relief, low birefringence (first order gray), length-fast orientation and a uniaxial interference figure. It was first thought to be zoisite but the association of skarn minerals indicated it was vesuvianite.

~~average~~ Small euhedral green garnets were scattered throughout the well-crystallized portions of these samples.

Average composition of B samples:

Vesuvianite	45%
Quartz	20%
Tourmaline	15%
Bornite	10%
Chalcite	3%
Calcite	7%
Garnet	3%
Sphalerite	1%

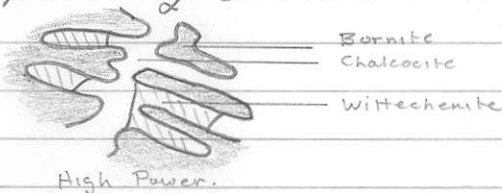
## Description of Metallic Minerals

1. Bornite This was the principal metallic mineral in the samples. Bornite was seen in large, fine grained masses with minor amounts of exsolved chalcocite, as striking myrmecitic intergrowths with chalcocite and wittichenite, and as exsolved grains in sphalerite and chalcopyrite. It was a purplish-brown colour which tarnished purplish-~~pt~~ blue rather quickly. Its low hardness and isotropism were conclusive factors in its identification.
  
2. Chalcocite This mineral was second to bornite in abundance in the samples. It was not readily visible in the hand specimens but was seen to occur with the bornite in very polished sections. Its commonest mode of texture was myrmecitic exsolution from bornite in fine, sub-oriented lamellae. It also occurred alone in fairly large grains, and as minute blebs and rounded inclusions in bornite, sphalerite and chalcopyrite. Its colour was white to bluish-white, tarnishing a light blue. It was very soft (B) and sectile. Anisotropism was moderate and etch tests were conclusive.
  
3. Sphalerite The sphalerite was of irregular distribution in the ore, and was not seen in all the sections. It was first thought to be tetrahedrite but further tests disproved this identification. The sphalerite was grey in colour and fairly hard with respect to the copper minerals above. It had a poor polish and a rather mottled surface. A reddish-brown internal reflection was seen under the carbon-arc light, and it was isotropic under reflected

light. Etch tests and microchem tests were conclusive. The sphalerite had a narrow, slightly convex boundary with the bornite and had inclusions of chalcocite in an emulsion texture.

4. Chalcopyrite Only one section showed chalcopyrite associated with bornite. The chalcopyrite seemed to be replacing the bornite and had inclusions of chalcocite. Veinlets of chalcopyrite cut into the bornite. No association of chalcocite and sphalerite were seen. Hardness, colour, isotropism and confirmatory etch tests identified this mineral.

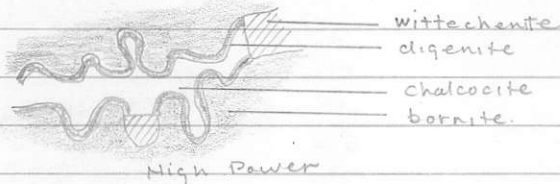
5. Wittechenite  $3\text{Cu}_2\text{S} \cdot \text{Bi}_2\text{S}_3$  This rare copper mineral was not readily identified. It occurred almost exclusively in the zones of fine myrmekitic bornite (chalcocite intergrowth) apparently replacing the small lamellae of chalcocite.



The wittechenite was pale creamy white, the same hardness as chalcocite and moderately anisotropic. The fine grain size and intimate association with other soft minerals prevented the removal of material for microchem tests. Etch tests were confirmatory.

6. Digenite  $4\text{Cu}_2\text{S} \cdot \text{CuS}$ . This mineral was observed in three of the sections. It surrounded bornite lamellae in a very thin rim texture, and was apparently replacing the <sup>bornite</sup> chalcocite. Its occurrence was restricted to bornite/chalcocite boundaries, and was seen almost exclusively in zones

of fine impurities intergrowth.

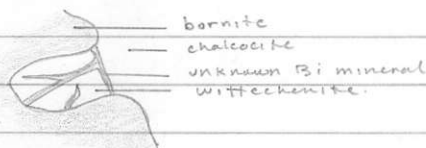


The texture of the digenite appeared granular but detailed examination was difficult due to the fine size of the digenite rim. It was a deep blue colour and isotropic, differing from chalcocite. Etch tests and hardness gave the same results as for chalcocite.

7. Pyrite A few isolated cubes and massive grains of pyrite were seen in non-steadily distributed fashion in the minerals. The small-sized grains had a high relief and a granular, poorly-polished surface. The hardness colour isotropism and roughly cubic shape positively identified the pyrite.

8. Covellite Thin films of covellite alteration were seen on many of the chalcocite grains in one section. The very high anisotropism with fiery orange colours, and the deep blue colour of the mineral were sufficient proof of identity.

9. Unidentified Bismuth Mineral This grey isotropic mineral was only seen under high power as tiny veinlets in, and rims around the Wittechenite. It term-



inated sharply at the wittechenite/chalcocite boundary and its intimate association with the wittechenite led to the belief it was a bismuth mineral. No etch tests or microchem tests could be made due to the small size of the grains.



## Physical, Chemical, and Optical Properties.

Mineral	Hardness	Colour	Etch Tests	Optics and Other Properties
Bornite $5\text{Cu}_2\text{S} \cdot \text{Fe}_2\text{S}_3$	B <sup>+</sup> less than chalcopyrite.	pinkish - brown to orange. tarnishes purple-violet	$\text{HNO}_3$ stains yellow - brown + eff. KCN - brown $\text{FeCl}_3$ - orange Cu microchem test.	Isotropic. Exsolution texture may be due to rapid cooling. Chalcocite is oriented in xlographic planes of host.
Chalcocite $\text{Cu}_2\text{S}$	B softer than bornite	bluish-white to white tarnishes blue.	$\text{HNO}_3$ stains blue. KCN - black $\text{FeCl}_3$ - blue Cu microchem test.	Moderate anisotropism - pinkish white to blue-grey. Blue streaks due to polish, tarnish. Replaced by wittechenite, digenite.
Sphalerite $\text{ZnS}$	C harder than chalcopyrite	grey	Zn microchem test positive.	Isotropic. Yellowish to reddish - brown int. refl. Poor polish.
Chalcopyrite $\text{Cu}_2\text{S} \cdot \text{Fe}_2\text{S}_3$	C	Brass - yellow	$\text{HNO}_3$ slow tarnish. Aq. regia - slow, brown differential stain Cu microchem test.	Weakly anisotropic. May exsolve from sphalerite. May be formed together with chalcocite by the disintegration of bornite.
Wittechenite $3\text{Cu}_2\text{S} \cdot \text{Bi}_2\text{S}_3$	B same as chalcocite.	creamy white	Aq. regia fumes stain. $\text{HNO}_3$ - slow eff.	Moderate anisotropism - light to dark grey. Smooth polish. Straight boundaries with chalcocite.
Digenite $4\text{Cu}_2\text{S} \cdot \text{CuS}$	B (chalcocite)	dark blue	same as chalcocite	Isotropic. Fine granular rim replacing chalcocite.
Pyrite $\text{FeS}_2$	F	yellowish - white	—	Poor polish. Granular surface. High relief. Isotropic. Cubic form.
Covellite $\text{CuS}$	B	Dark blue	—	Very high anisotropism - fiery orange - brown

## Abundance of Primary Metallic Minerals

An average of 10 cross-mounted polished sections, 6 plasticine mounts and 13 hand sampled was taken, giving the following composition:

Bornite	56%
Chalcocite	22%
Sphalerite	15%
Chalcopyrite	3%
Wittichenite	2%
Digenite	1%
Pyrite	1%
Bismuth mineral	<1%

Only one alteration mineral, covellite, was seen in polished sections, but other supergene copper minerals were seen in the hand specimens. The minerals formed by alteration were:

Covellite  
Malachite  
Azurite  
Limonitic copper salt(?)

No estimate of abundance was made.

## Textures and Paragenetic Sequence

The earliest-formed minerals were non-metallic minerals: garnet and vesuvianite, then tourmaline and quartz in that order. Borneite veins cut through the vesuvianite and quartz, and massive borneite surrounds the garnet and tourmaline in hand specimens. Calcite is either earlier than the borneite or contemporaneous. One hand specimen indicates simultaneous crystallization by the texture, but the calcite later was recrystallized.

The borneite / chalcocite boundary relation is almost always a symplectitic intergrowth. This could be a pseudotectite texture from exsolution after simultaneous deposition, or a replacement texture. Some boundary relations between borneite / chalcocite seem to indicate replacement of the chalcocite by the finger-like projections of the borneite into the chalcocite, but the rounded nature of the lamellae and the equivalent lamellae of chalcocite into borneite cause some uncertainty about the replacement relationships. The boundaries between the two minerals are essentially smooth curves. Inclusions of chalcocite in the borneite are randomly distributed, and of a rounded, oval or sub-oval shape indicating contemporaneous deposition and/or exsolution, rather than replacement. A certain degree of replacement is believed to have occurred, i.e. borneite replacing chalcocite or chalcocite replacing borneite, but contemporaneous deposition of the two minerals predominates.

The Wittichenite appears to have selectively replaced the chalcocite in the zones of symplectitic intergrowth. It always

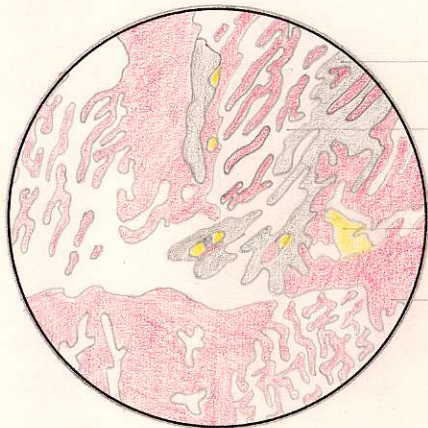
forms a smooth boundary with the chalcocite, but with a convex surface towards the chalcocite or, as in some cases, rounded projections into it. Some small exsolved inclusions of chalcocite in bornite show partial replacement by Wittenhite, with the above boundary relationships. The Wittenhite has not replaced the bornite in any samples, although it is in direct contact with it.

One interesting feature of the Wittenhite is the unidentified mineral seen associated with it in most of its occurrences. The mineral is believed to be a bismuth mineral because of its association with Wittenhite alone. In section 4, it appears primary, <sup>because</sup> (since) it forms symplectitic intergrowths with bornite and chalcocite, and has inclusions of bornite, chalcocite and Wittenhite. As it was with the bornite/chalcocite relationship, this may be either a primary pseudotectic texture or a replacement texture. However, the bismuth mineral shows definite replacement texture in the Wittenhite, with long, vein-like projections cutting the Wittenhite but terminating at the Wittenhite/chalcocite boundary. It also shows a rim texture enclosing the Wittenhite. The "primary texture" exhibited in the symplectitic bismuth intergrowth may be a complete replacement of the Wittenhite which had itself replaced chalcocite, or it may be a replacement of the chalcocite and bornite, along with the Wittenhite.

In several sections, a narrow dark blue rim of digenite was seen at the boundary of the bornite/chalcocite. It was first thought to be simply

an alteration of the chalcocite to the isotropic form, but closer examination showed isolated lamellae of bornite in chalcocite which had partly or wholly altered to digenite. This alteration replaced the chalcocite to the extent that the rims of digenite on opposite sides of narrow chalcocite lamellae will meet. It was noted that the digenite alteration occurred exclusively at the boundary between bornite and chalcocite. Sometimes the bornite was replaced (wholly altered lamellae) and sometimes the chalcocite was replaced (rims of digenite inside enclosed lamellae of chalcocite). Digenite did not occur at the chalcocite / wittichenite boundary, the wittichenite apparently replacing the digenite in the process of replacing the chalcocite.

Sphalerite was observed to be in contact with 5 other minerals: bornite, chalcocite, wittichenite, digenite and pyrite. In general it had rounded, convex or smooth boundaries with these minerals. In polished section 8, sphalerite is seen to cut across chalcocite, bornite and digenite, and project into wittichenite, indicating replacement of all these minerals. Section 8 also showed inclusions of chalcocite in the sphalerite in an emulsion texture. Similar textures were seen in other sections. This resolved chalcocite was obviously of a later stage than the primary chalcocite. Section 6 shows sphalerite replacing bornite and also resolving bornite in a small rounded inclusion. Like the chalcocite mentioned above, this bornite is of a later stage than the primary bornite.



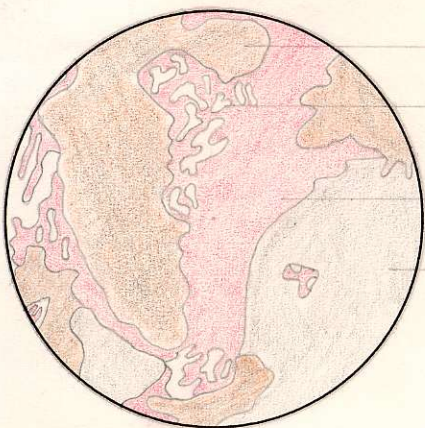
Bismuth mineral - grey

Chalcocite - white

Wittechenite - yellow

Bornite - red

Polished Section 4  
Med. Power



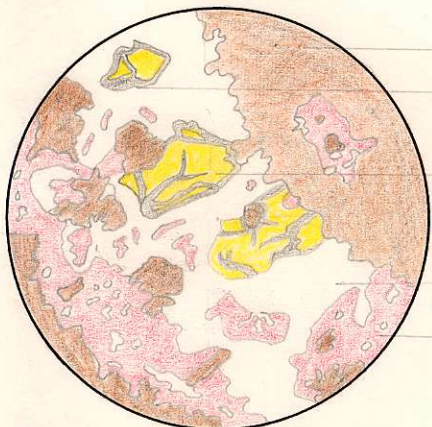
Gangue (garnet) - brown

Chalcocite - white

Bornite - red

Sphalerite - grey

Polished Section 6  
Med. Power



Gangue (garnet) - brown

Bismuth mineral - grey

Wittechenite - yellow

Chalcocite - white

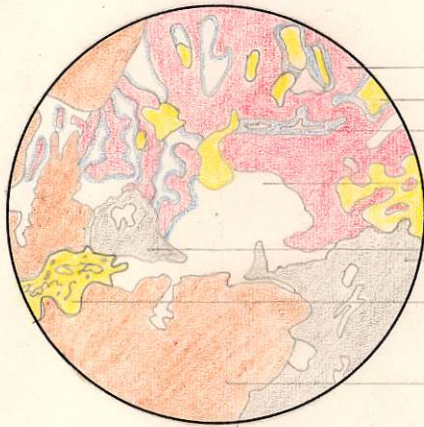
Bornite - red

Polished Section 7  
High Power

An interesting texture seen in this small inclusion, and in similar inclusions elsewhere, was the bornite altering to chalcocite in two small areas.

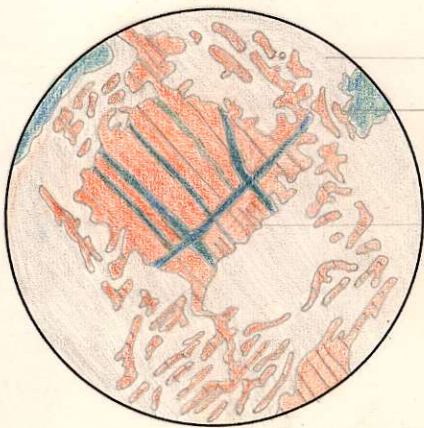
Chalcopyrite was seen only in one polished section, mount III, and boundary relations with only three minerals: bornite, chalcocite and pyrite, were observed. Hence the true paragenetic relationships between chalcopyrite and all the minerals later than chalcocite, with the exception of pyrite, could not be determined. The chalcopyrite had fairly regular boundaries with the bornite and chalcocite, except for a few prominent remnants of chalcocite cutting those two minerals. Replacement was evident by examination under high power which showed a serrated boundary between chalcopyrite and bornite with the cusps into the bornite. Convex boundaries and small projections were seen from chalcopyrite to bornite and from chalcopyrite to chalcocite. Very small inclusions of chalcocite were seen in a roughly defined zone within the chalcopyrite near the boundary with bornite. This inclusion texture was evidence of a late stage of chalcocite resolved from the chalcopyrite.

The last mineral to crystallize was pyrite. It was seen in only three sections, and then as minute, isolated grains. In section B, it was seen to replace chalcopyrite, chalcocite, bornite and sphalerite and in mount III it replaced chalcopyrite. The pyrite commonly occurred in cubic grains standing in high relief and cutting across the other metallic minerals.



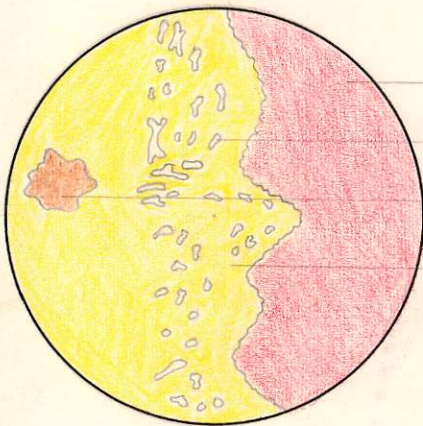
- Bornite - red
- Witchechenite - yellow
- Digenite - blue
- Chalcocite - white
- Sphalerite - grey
- Pyrite - dark yellow
- Gangue (garnet) - brown

Polished Section 8  
High Power



- Bornite - grey
- Gangue (quartz) - green
- Covellite - red

Plasticine Mount II  
X-Niols, High Power



- Bornite - red
- Chalcocite - white
- Gangue (garnet) - brown
- Chalcopyrite - yellow

Plasticine Mount III  
High Power



Covellite was seen as an alteration film on chalcocite in plastic mount II. The alteration film was too thick to be removed by buffing. The section was unique in that it was the only chalcocite which showed cleavage. A large grain of chalcocite was surrounded by smaller graphitic or myrmecitic grains of chalcocite, all of which were optically continuous with the large grain and also showed continuous cleavage with it. The replacing mineral was bornite. Covellite had selectively replaced the chalcocite. Quartz gangue penetrated the chalcocite along cleavage planes and also cut across the cleavage. It appears that silicification took place before bornite replacement since the quartz terminated abruptly at the chalcocite/bornite boundary.

### Nature of Mineral Deposit

The gangue minerals present were typical of a steam deposit: garnet, tourmaline, vesuvianite and calcite. Some minor silicification was present also. The metallic minerals were deposited later and showed replacement textures indicative of metasomatic replacement. The common sulfides present; bornite, sphealerite, chalcocite and chalcopyrite, were hypogene minerals in the mesothermal temperature range. All evidence indicates a contact metasomatic or steam-type deposit.

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RMT  
Nov. 30/63

# Paragenetic Sequence

