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DOLMAGE CAMPBELL & ASSOCIATES (1975) LTD.
CONSULTING GEOLOGICAL & MINING ENGINEERS
1000-1055 WEST HASTINGS STREET
VANCOUVER, CANADA V6E 2E9

Georgia Resources Inc.

Geological Report on

TRAPPER LAKE MOLYBDENUM PROPERTY

Northwestern British Columbia

NTS, 104K/7

September 25, 1980

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Consultants
Vancouver, Canada

report taken from VSE, Glory Exploration Ltd

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ILLUSTRATIONS

DOLMAGE CAMPBELL & ASSOCIATES (1975) LTD.
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SUMMARY

The Trapper Lake property, located in northwestern British Columbia, is underlain by molybdenum-bearing quartz monzonites or alaskites which are typical of commercial molybdenum deposits in the Cordilleran belt of North America. The molybdenum occurs as molybdenite in fractures associated with quartz, and is distributed non-uniformly over an L-shaped zone some 2700 metres in length. One of three holes drilled near the north end of the area in 1972 intersected a 15.2 metre zone averaging 0.18% MoS_2.

The mineralized areas require additional prospecting and geological mapping. Diamond drilling is needed both to test the mineralized areas and to explore extensions of such areas beneath boulder gravel outwash which covers much of the valley floor.

The recommended exploration program is summarized below:

Diamond drilling (9 holes)	$195,000.
Helicopter, transportation	27,000.
Engineering, mapping, camp costs	48,000.
	$270,000.

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INTRODUCTION

In August, 1980, the firm of Dolmage Campbell & Associates (1975) Ltd. was requested by Mr. Seamus Young of Georgia Resources Inc. to write a geological report on the Trapper Lake molybdenum property of north-western British Columbia. One of the authors (C.R. Saunders) supervised an exploration program on the property in the summer of 1972 and personally mapped some of the molybdenum bearing zones. The present report is based on data existing in Dolmage Campbell files (see references) evaluated in the light of current and projected molybdenum prices.

LOCATION AND ACCESS: (58° 18'N, 132° 39'W)

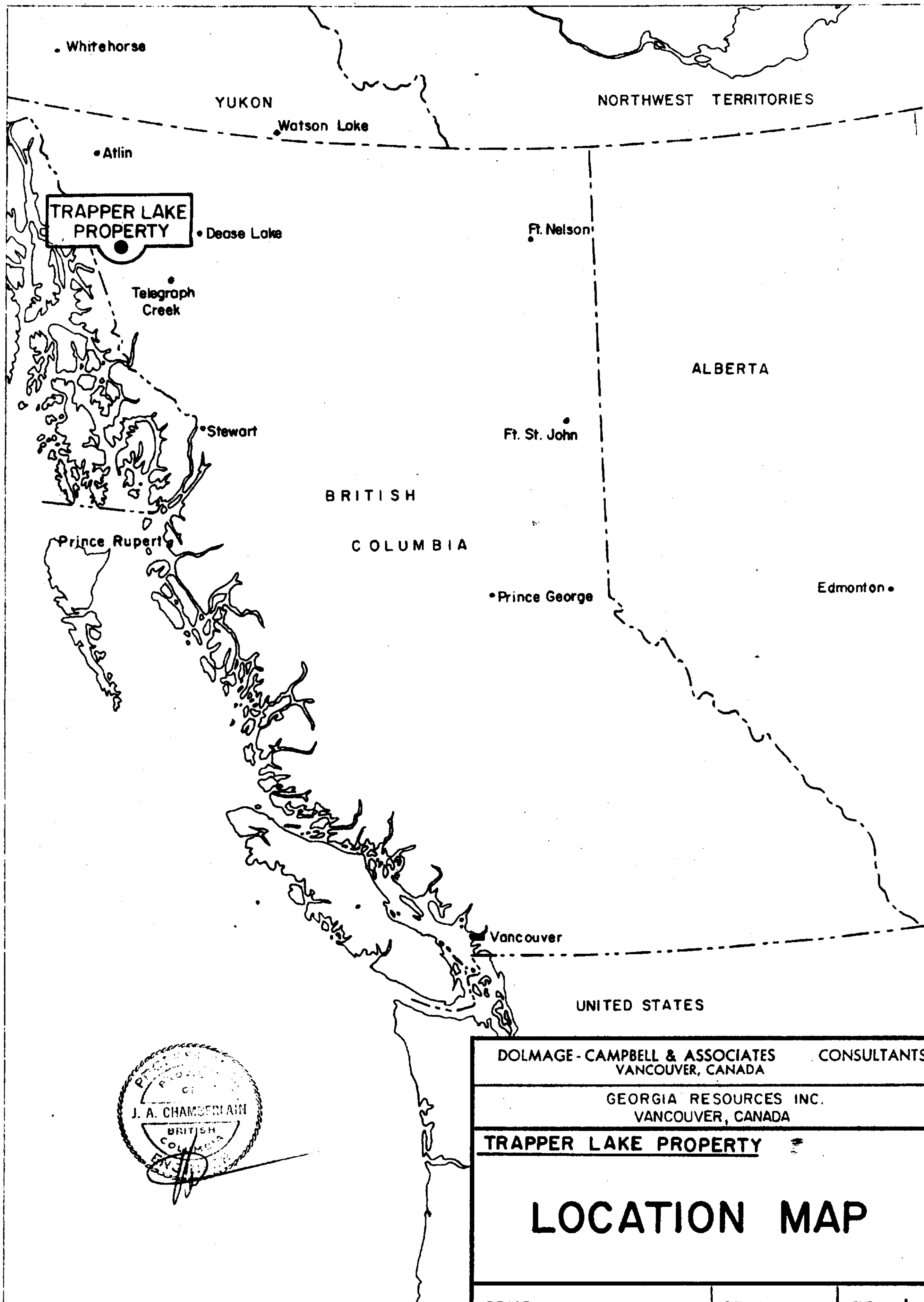
The molybdenite property of Georgia Resources Inc. is located in the northwest corner of British Columbia about 145 kilometres south and east of Atlin and 100 kilometres northwest of Telegraph Creek, (Figure 1). The property is equidistant (290 kilometres) from Whitehorse, Y.T., to the north-northwest, from Stewart to the south-southeast and from Watson Lake to the northeast. The city of Juneau, Alaska, lies 105 kilometres due west of the property but access to it is impractical except by air.

The property is accessible only by air, either by helicopter directly to the property or by float plane to Trapper Lake which lies 16 kms north of the property. A smaller lake, 6 kms north of the property, can be used by small float planes from early July until freeze-up, but use of this lake is not recommended for larger float planes with heavy loads.

During the 1972 exploration program, the camp was supplied by air from Atlin. It is possible that some economies could have been effected by supplying from Dease Lake, but on that occasion, good communications through Atlin had already been set up and supplies were more readily available at short notice through Atlin than through Dease Lake. The closest road is the Dease Lake-Telegraph Creek road about 100 kilometres to the southeast. The Cassiar-Stewart Highway, at Dease Lake, lies 150 kilometres due east of the property.

PROPERTY

The Trapper Lake property of Georgia Resources Inc. comprises six mineral claims, as follows:



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GEORGIA RESOURCES INC.
VANCOUVER, CANADA

TRAPPER LAKE PROPERTY

LOCATION MAP

SCALE: 1 7,603,200	SEPT. 1980	FIG. 1

TIM No. 1	(16 units)
TIM No. 2	(12 units)
TIM No. 3	(20 units)
TIM No. 4	(20 units)
LINDA	(8 units)
LEAH	(8 units)
	(84 units) (2100 ha.)

All are owned by Georgia Resources Inc.

The location of the claim groups is shown in Figure 2.

HISTORY

Because of difficult access the northwestern corner of British Columbia has been only sparsely explored to date. Placer gold deposits were exploited east of Atlin, and the Polaris Taku copper-lead-zinc mine on the Taku River produced until the late 1950's, but it was not until after World War II and the beginning of construction of the Cassiar-Stewart Highway that modern prospecting began to penetrate the area along the east flank of the Coast Range inland from the Alaska Panhandle. This resulted in the discovery and development to production of the Granduc (copper), Cassiar (asbestos) and Nu-Energy Resources (gold) deposits. Still in the development stage are Adanac (molybdenum), Schaft Creek (copper-molybdenum) and a number of other deposits.

The Trapper Lake molybdenum showings were discovered and staked by Southwest Potash Corporation, a subsidiary of American Metals-Climax in 1962. The property was restaked in 1970 and optioned to Plateau Minerals and Industries Ltd.

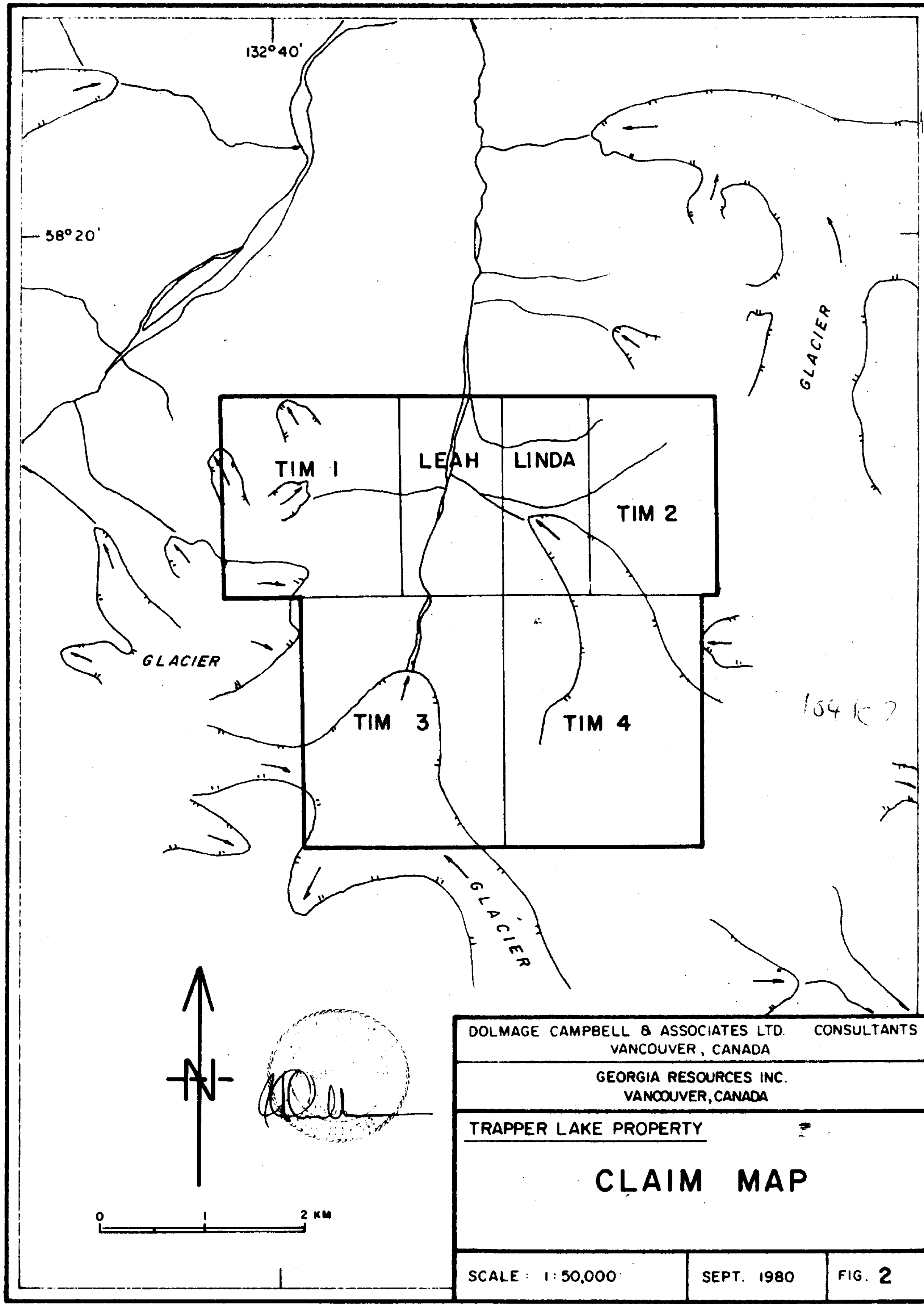
The first physical work on the property was done in 1972 and consisted of the detailed geological mapping of Zone A (west) and Zone C (east), the reconnaissance geological mapping of zones B (central), C and D (south), and the drilling of a total of three holes on zones A and C (317 metres).

The claims were allowed to lapse and the ground was re-staked in 1980 by Georgia Resources Ltd.

REFERENCES

The following publications were reviewed in the preparation of this report:

"Geology and Mineral Deposits of Tulsequah map-area, British Columbia", 1971; Memoir 362, Geological Survey of Canada; J.G. Souther.



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GEORGIA RESOURCES INC.
VANCOUVER, CANADA

TRAPPER LAKE PROPERTY

CLAIM MAP

SCALE: 1:50,000	SEPT. 1980	FIG. 2

"Report on Trapper Lake Molybdenum Property, Northwestern B.C.", November 19, 1971; Dolmage Campbell & Associates Ltd.

"Summary Report on the Trapper Lake Property, Northwestern B.C.", Jan. 15, 1973; Dolmage Campbell & Associates Ltd.

GEOLOGICAL SETTING

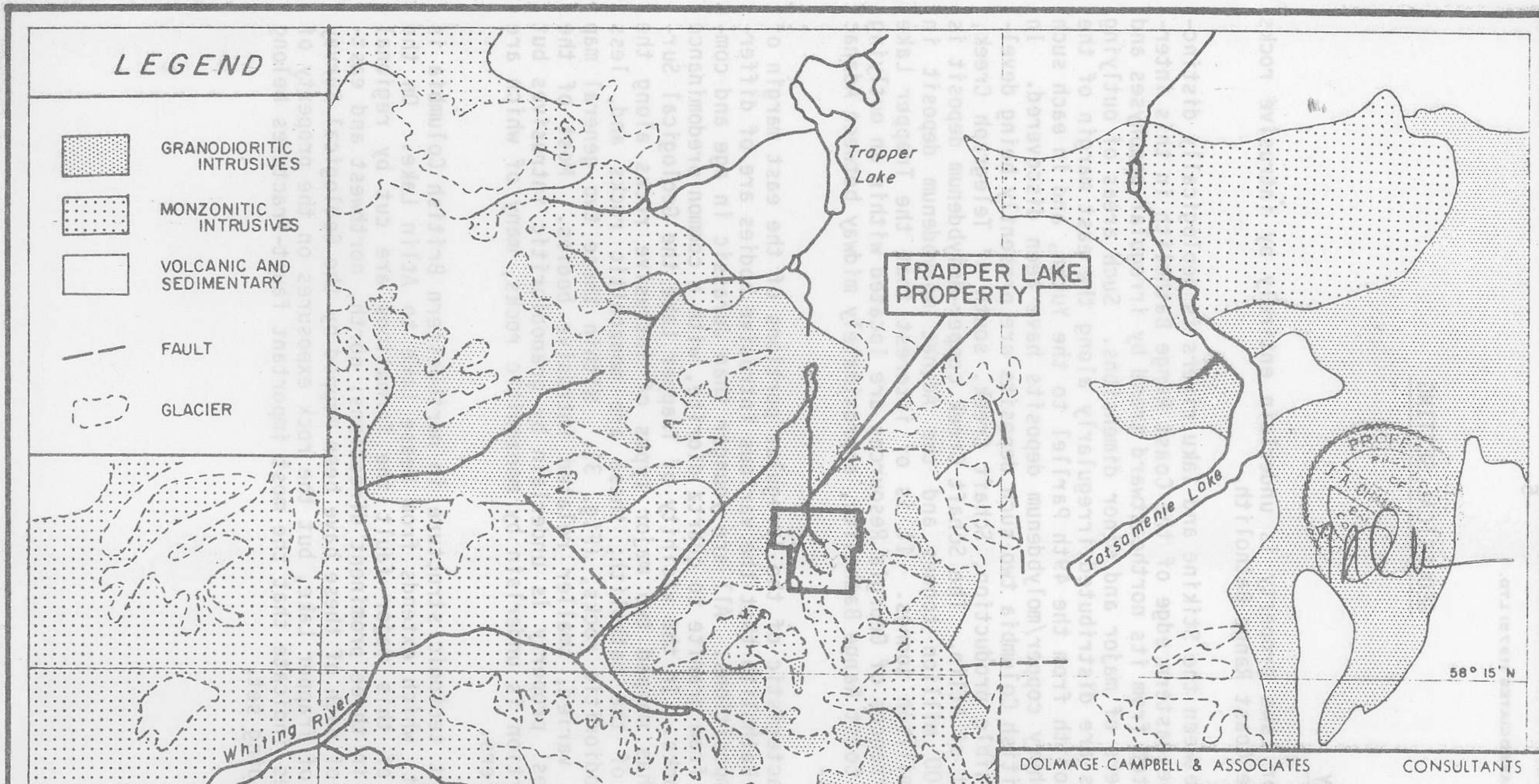
REGIONAL GEOLOGY

The Trapper Lake property is underlain entirely by intrusive rocks belonging to the Coast Range Batholith.

The area between the Stikine and Taku rivers is geologically distinctive in that the eastern edge of the Coast Range Batholith in this interval is interrupted from its northeastward trend by irregular apophyses and intrusive outliers of major and minor dimensions. Such areas of outlying intrusive bodies are distributed irregularly along the east margin of the Coat Range Batholith from the 49th Parllel to the Yukon, and in each such area major porphyry copper/molybdenum deposits have been discovered. In northwestern British Columbia two such deposits are presently being developed toward possible production, Schaft Creek, south of Telegraph Creek, and Adanac, east of Atlin. The Schaft Creek copper-molybdenum deposit is in excess of 300 million tonnes and the Adanac molybdenum deposit in excess of 100 million tonnes. It is of interest that the Trapper Lake molybdenum occurrences of Georgia Resources are located within an outlying apophyse of the Coast Range Batholith, approximately midway between Adanac and Schaft Creek.

It is characteristic of the irregular portions of the east margin of the Coast Range Batholith that the various intrusive bodies are of different ages and compositions. All are younger than Jurassic in age and compositions range from granite to quartz diorite, with a common predominance of monzonitic rocks. In the vicinity of Trapper Lake the Geological Survey of Canada has mapped two major types of intrusive rocks along the eastern margin of the batholith; these are monzonitic rocks and, less extensive, granodioritic rocks, (Fig. 3). Within these two general map units are many, varied, smaller related intrusive bodies. Most of the Georgia Resources' property is underlain by granodioritic intrusives but the southern portion is underlain by monzonitic rocks, many of which are covered by glaciers.

The dominant tectonic structure in northwestern British Columbia is the Nahlin Fault, which extends from Dease Lake to Atlin Lake. On the southwest side of this major fault the formations are cut by regional faults trending in three prominent directions; north, northwest and eastward. Examples of each of these have been mapped by the Geological Survey in the vicinity of Trapper Lake, but the rock exposures on the property of Georgia Resources indicate that the most important fault-fractures belong to the east-trending set.


LEGEND
GRANODIORITIC INTRUSIVES
MONZONITIC INTRUSIVES
VOLCANIC AND SEDIMENTARY
FAULT
GLACIER
Trapper Lake
TRAPPER LAKE PROPERTY
Totsamenie Lake
Whiting River
B.C.
ALASKA
58° 15' N
133° W
132° 30' W
DOLMAGE CAMPBELL & ASSOCIATES
CONSULTANTS
VANCOUVER, CANADA
GEORGIA RESOURCES INC.
VANCOUVER, CANADA
TRAPPER LAKE PROPERTY
REGIONAL GEOLOGY
SCALE 1 : 253, 440
SEPT 1980
FIG 3

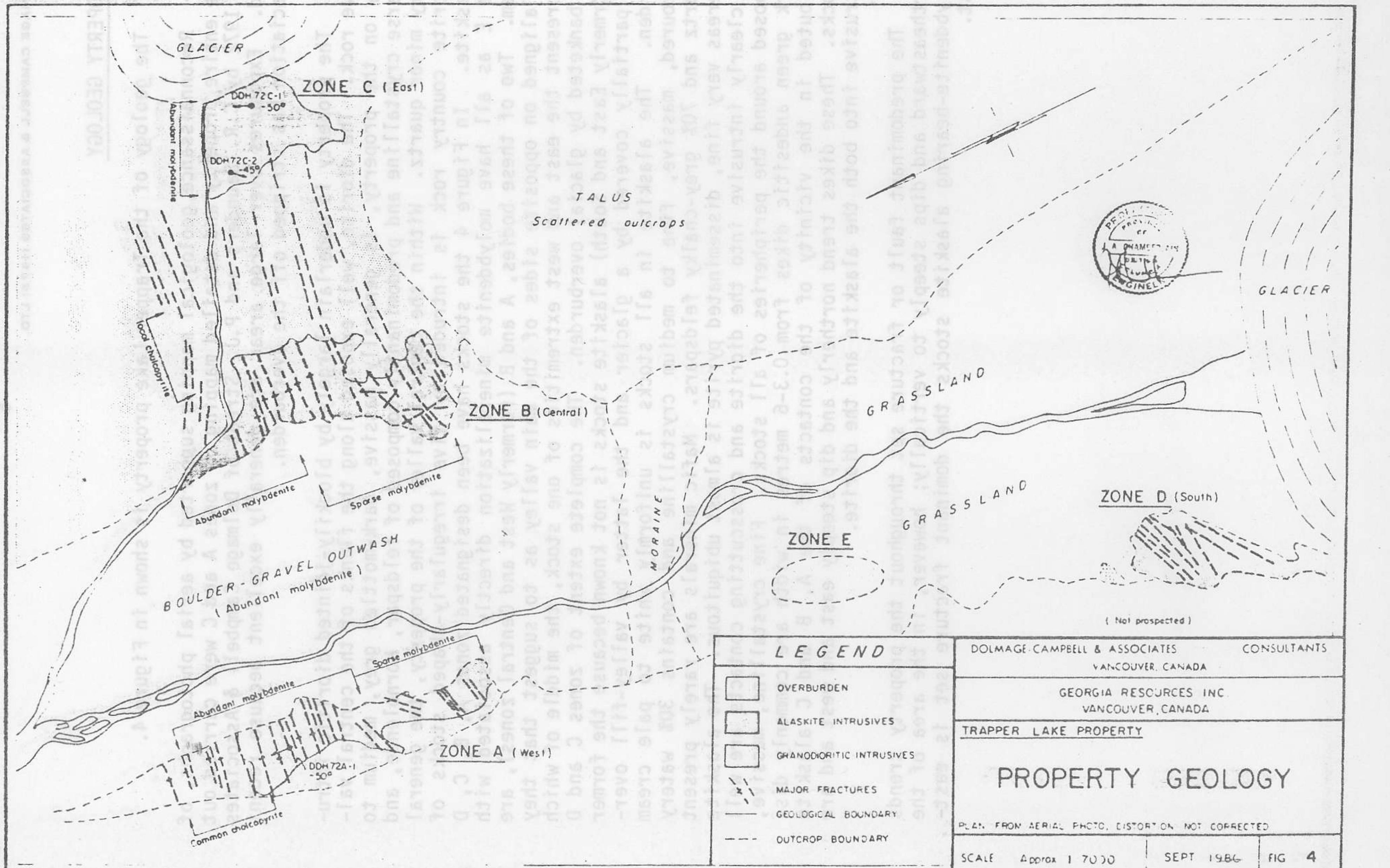
PROPERTY GEOLOGY

The geology of the Trapper Lake property is shown in Figure 4.

Reconnaissance geological mapping supported by aerial photographs, of the entire property and detailed mapping of zones A and C were carried out in 1972 by C.R. Saunders and P.J. Street of Dolmage Campbell & Associates Ltd. Exposures over large areas are generally excellent because recent glaciation has stripped off the overburden.

The property is underlain largely by blockily-jointed dioritic intrusive rock. The diorite, well exposed along the flanks of the central valley on the property, is generally massive, dark-mottled grey, medium to coarse crystalline and predominantly composed of feldspar, hornblende, and very minor quartz. Within the central valley of the property, the general diorite country rock is intruded by five irregularly-shaped stocks of alaskite. In Figure 4 the stocks have been designated zones A, B, C, D and E as all have molybdenite mineralization directly associated with them. Two of these bodies, A and B (formerly West and Central zones), are so aligned on opposite sides of the main valley as to suggest that they represent the east and west extremities of one stock, the middle of which is banketed by glacial overburden. The complete extent of zones C and D (formerly East and South) alaskite stocks is not known because the former is partially covered by a glacier and the latter by valley-fill overburden. The alaskite in all stocks is uniformly white to pale cream coloured, massive, fine to medium crystalline and contains 30% watary quartz and 70% grey-chalky feldspars. Mafic minerals are rarely present whereas very fine, disseminated pyrite is almost ubiquitous. The alaskite is clearly intrusive into the diorite and crosscutting contacts are well exposed around the peripheries of all stocks. Fine crystalline, massive, dark green andesitic dikes from 0.3-6 metres in width are commonly distributed in the vicinity of the contacts of the A, B and C alaskite stocks. These dikes trend northerly and dip steeply east and west and are intrusive into both the alaskite and the diorite.

The predominant fault or fracture set throughout the property trends northeastward and dips steeply to vertically; however, in the area of the molybdenite-bearing alaskite stocks the dominant fracture set is east-west.


GLACIER
ZONE C (East)
DDH 72C-1
-50°
DDH 72C-2
-45°
Abundant molybdenite
Local chalcopyrite
TALUS
Scattered outcrops
ZONE B (Central)
Abundant molybdenite
Sparse molybdenite
BOULDER - GRAVEL OUTWASH
(Abundant molybdenite)
MORAINE
GRASSLAND
GRASSLAND
GLACIER
ZONE D (South)
ZONE E
(Not prospected)
Sparse molybdenite
Abundant molybdenite
ZONE A (West)
DDH 72A-1
-50°
Common chalcopyrite
LEGEND
OVERBURDEN
ALASKITE INTRUSIVES
GRANODIORITIC INTRUSIVES
MAJOR FRACTURES
GEOLOGICAL BOUNDARY
OUTCROP BOUNDARY
DOLMAGE-CAMPBELL & ASSOCIATES
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GEORGIA RESOURCES INC.
VANCOUVER, CANADA
TRAPPER LAKE PROPERTY
PROPERTY GEOLOGY
PLAN FROM AERIAL PHOTO, DISTORTION NOT CORRECTED
SCALE Approx 1:7000
SEPT 1986
FIG 4

ECONOMIC GEOLOGY

Molybdenite is the mineral of predominant economic interest on the Trapper Lake property. It is exposed in outcrops over wide areas in the alaskite bodies and locally into the adjacent diorite. The molybdenite occurs as disseminations in the rock, as thin films and blebs on fracture faces and with quartz and pyrite in veinlets. In most of the surface exposures the molybdenite has undergone some leaching. The concentrations of molybdenite are variable in any outcrop from trace to locally very abundant, thus making estimates of possible grade very difficult.

Chalcopyrite occurs in local abundance in the outcrop of diorite in Zone A, mostly along east-west, fine fractures.

In 1972 a program was undertaken to map the best molybdenite showings in detail and to drill a number of preliminary diamond drill holes into the better surface showings of zones A and C (Figure 4).

Description of the principal mineralized exposures that were mapped in 1972 are given below. The showings were mapped in the field on 100-scale (1:1200) and these field maps are available in the files of Dolmage Campbell & Associates (1975) Ltd.

Three core holes were drilled on the Trapper Lake Property during 1972. The total amount drilled was 317 metres. One hole was drilled on Zone A and two holes were drilled on Zone C. No drilling was done on zones B, D and E. Results of the drilling are discussed under the appropriate zone.

ZONE A

Geology: A central body of alaskite, elongate in a north-south direction, is flanked on the west and north by diorite, and is covered on the east by glacial and glacio-fluvial overburden. Alaskite is in contact with quartz monzonite on the south margin, but as outcrop is not abundant here, this change of rock-type may be of local extent only.

The alaskite consists of quartz (10-50%), feldspar (predominantly orthoclase, but with local abundant plagioclase), and minor amounts of hornblende and biotite. It varies from place to place in relative content of quartz, orthoclase or plagioclase. It is generally medium- to coarsely-crystalline, and has no directional structure such as gneissic foliation. It shows no particularly strong alteration near the contact with the diorite, but has undergone slight to moderate chloritization and/or kaolinization throughout. Mineralized fractures are almost invariably chloritized to some extent.

In the diorite, plagioclase feldspar predominates, quartz is generally less than 5%, and mafic minerals amount to as much as 30%. The diorite is commonly well foliated but not gneissic, i.e. feldspars and mafic minerals show planar alignment, but without marked segregation into light and dark layers. Adjacent to the contact at the northern end of the alaskite, the diorite is highly altered in a rusty-weathering zone up to 7-9 metres wide, in which the original texture of the diorite is blurred or obliterated. Elsewhere along the contact, the diorite is very little altered.

Surface Mineralization: Both the alaskite and the diorite are cut by intersecting sets of fractures, mostly vertical or steeply-dipping, that strike predominantly east-west to about 290°, and from north-south to as much as 030°. There is a lesser set at about 050°, dipping at 65° to the southeast. The fractures that strike at 270° to 290° dip individually either north or south, at angles ranging from 25° to 85°. Spacing on all sets of joints varies greatly from place to place, but is consistently closer (from a few centimetres up to 0.6 metres) within 100 to 125 metres on either side of the northern contact than it is farther away from that contact. Close fracturing is also associated with several prominent steeply-dipping shear zones that strike approximately southwest, 310° and 340°.

Medium- to coarsely-crystalline molybdenite is found in quartz veins, in fractures and, more rarely, disseminated within the unfractured rock.

In quartz veins, the molybdenite generally forms discontinuous layers close to the walls of the vein, as much as 1 cm thick in a few places but more commonly much thinner. In a few instances, molybdenito forms layers within the quartz vein as well as on its margins, and in other instances, the molbdenite is intimately mixed with finely- to coarsely-crystallized pyrite. In fractures, it forms coarse rosettes and blebs, or finer flecks scattered over the fracture surface, and where movement has occurred on fractures, the molybdenite is thoroughly smeared over the slickensided surface. The molybdenite found in fractures is rarely as much as 2 mm thick, most commonly much thinner. As disseminated material in non-fractured rock, it occurs only sparsely, and, from core assays it is evident that it makes no appreciable difference to the grade. Coarse flecks up to 5 mm are seen, but most are 1-2 mm or smaller.

Molybdenite on veins and fractures is most abundant in the alaskite within 100-125 metres of the diorite contact at the north end of the A Zone, but an appreciable amount also occurs in the diorite close to the same contact. Molybdenite in fractures is less common in the diorite than in the alaskite, but most quartz veins within 100 metres of the contact carry molybdenite mineralization. Disseminated molybdenite in the diorite is rare.

All quartz veins, both in alaskite and diorite, contain abundant fine to very coarsely-crystallized pyrite, either sugary-textured or in euhedral crystals up to 4-5 cms across. The sugary pyrite may be mixed with molybdenite. Finely-crystallized pryite is disseminated throughout the alaskite, but very sparsely, and nowhere exceeds 1% of the rock. The diorite has only traces of disseminated pyrite.

Diamond Drilling: Diamond drill hole 72-A1 was drilled north at minus 50° to a depth of 119.5 metres and was located near the north contact of the Zone A alaskite body to intersect as many as possible of the major, significantly-mineralized fracture-sets, in this, the best surface showing on the property. The entire core from DDH 72-A1 was shipped for assay. In addition, four sludge samples were assayed, each from sections approximately 7.5 metres in length.

The hole remained in alaskite throughout is length. The rock is fine to coarsely crystalline, massive, white to pale pink, with weak to moderate chloritic or kaolinitic alteration locally, especially along fractures. Quartz veins up to 1 cm in width which occur intermittently through the section are associated with molybdenite. Most 3-metre sections assayed less than 0.05% MoS_2 but one 15.2 metre section from 93.6 to 108.8 metres averaged 0.18% MoS_2.

ZONE B

Geology: Alaskite and quartz monzonite are exposed in the B Zone over an area about 525 metres long from north to south, from the north bank of the creek that drains the C Zone glacier, to the lateral moraine left by that glacier when it was much farther advanced than it is at present. The Zone is 240 to 270 metres wide, but is covered on the west flank by glacial overburden, and it could therefore be continuous with the A Zone on the west side of the valley.

Surface Mineralization: Molybdenite in the B Zone occurs in the same manner as in the A Zone, but in spite of occasional conspicuous molybdenite rosettes and blebs on exposed fracture surfaces, it is much less abundant. In particular, quartz veins are much more poerly developed than in the A Zone. The zone is cut by several andesite dykes, up to 1.5 metre wide, striking approximately north-south and dipping steeply, but no mineralization is associated with them. A major rusty-weathering shear zone, with appreciable calcite, striking 025°, cuts the diorite close to the east margin of the alaskite. It appears to carry no molybdenite. The principal fracture directions in the alaskite strike east-west and north-south, and most are steeply- to vertically-dipping.

ZONE C

Geology: Where exposed, the alaskite in the C Zone forms an L-shaped body, the angle of the L pointing roughly northeast, and each arm about 350 metres long. The east-west arm is relatively narrow, averaging 100 metres in width. The north-south arm is over 125 metres wide where exposed, and its total extent is not seen because it is covered by tho toe of a glacier. There are numerous extensive outcrop areas up to 100-125 metres in length, but outcrop is intermittent over most of the alaskite body because of covering by great volumes of glacially-deposited debris.

The alaskite is similar in composition to that in the A Zone, generally fine- to medium-crystalline, and mostly showing weak chloritic and kaolinitic alteration. Most of it is white or groenish-white, but shows a faint pinkish tint in a few places, particularly in the southeast portion of the zone. There seems to be no correlation of this colour variation with intensity of mineralization. A white, sugary-textured, finely-crystalline alaskite, commonly well-mineralized where it is seen in float on the valley floor, was not found in place in the C Zone.

Contacts of alaskite with the diorite country-rock are clean, and alteration of the alaskite is no more intense at the contacts than elsewhere. A few narrow inclusions and discontinuous dykes of alaskite are found within the diorite close to the contact.

The diorite is identical to the diorite in A Zone; however, in C Zone the diorite in contact with alaskite is nowhere as strongly altered as it is at the northern contact with the alaskite in A Zone.

The diorite shows a moderate degree of foliation, but the alaskite is invariably massive.

Surface Mineralization: The directions of strike of joint sets correspond closely to those in A Zone, i.e. most commonly east-west, north-south and north-northeast, but the dip is more variable ranging from vertical to as little as 20°. For all strike directions, the steeper dips vary either side of the vertical, but a distinctive feature of tho C Zone structure is the abundance of shallow (20-30°), north-northwest- and northeast-dipping fractures with strikes approximately east-west and north-northeast. These fracture sets are noteworthy because they commonly form dip-slopes and in many places carry conspicuous molybdenite mioeralization. Spacing of joints varies from a few centimetres to about one metre, but is not consistently close.

Disseminated molybdenite is rare in the C Zone. Coarse flecks, blebs and rosettes of molybdenite are conspicuous throughout the zone, in large part because they so commonly occur on exposed fracture surfaces with a shallow dip. All directions of fracture carry some molybdenite, but it is most abundant in fractures with approximate east-west strike. There are small areas of local concentration, for example, in the creek-bed at the northeast corner of the zone, in two outcrops roughly southwest of DH 72-C2, and near the glacier at the southeast end of the zone. These may indicate that mineralization is stronger close to the diorite contact than elsewhere.

Mineralization on fractures is nowhere more than about 3 mm thick, and is generally much thinner. In places, the molybdenite is associated with vein quartz, but the latter is invariably very narrow, veins over 1 cm being rare. In general, vein quartz in the C Zone is intermittent and sparse, and does not compare with the relatively well-developed system of east-west striking quartz veins in the A Zone. In a few places, molybdenite is smeared over fracture surfaces on which slight movement has taken place.

Pyrite is much more sparse in C Zone than in A Zone. There is very little finely-crystallized disseminated pyrite, and only a few fractures are seen that carry more coarsely-crystallized pyrite. The finely-mixed pyrite and molybdenite that occurs in some quartz veins in the A Zone is not seen in the C Zone. No chalcopyrite is seen in or close to the C Zone alaskite, but traces in diorite boulders on the valley floor indicate that it is present somewhere in the B or C Zones.

Diamond Drilling: Two holes were drilled on the C Zone as follows:

No.	Length	Dip	Azimuth
72-C1	(64.6 m)	-50°	020°
72-C2	(133.2 m)	-45°	225°

Both holes remained in alaskite throughout their lengths. In general, the rock appears to be similar to that intersected in hole 72-A1 but with less quartz veining. Unfortunately, only a few sections were assayed from these holes: two from 72-C1 and six from 72-C2. The two best assays were from hole 72-C2. A 3.3 metre interval from 20.4 to 23.7 metres assayed 0.55% MoS_2 and a 1.8 metre interval from 53.6 to 55.4 metres assayed 0.86% MoS_2.

ZONES D, E

These zones are located on Figure 4. Zone D is exposed near the toe of the south glacier and Zone E lies roughly half way between zones D and A. Both zones D and E are underlain by alaskite containing sporadic molybdenite in fractures and disseminations. Neither one has been prospected in detail.

DISCUSSION

The results of the 1972 preliminary drilling program at Trapper Lake were not encouraging at the time. However, during the last few years, the requirement for molybdenum in a vastly wide range of alloying uses has placed heavy demands on producers. The price of molybdenum in 1972 was $1.75 per pound versus a current price in excess of $10.00 per pound - an increase in the order of 6 times. As a comparison, during the same period, the price of copper increased from 50 cents to 90 cents per pound - an increase of less than 2 times.

The best zone assayed in the three holes drilled in 1972 was a 15.2 metre intersection averaging 0.18% MoS_2. While not spectacular, this grade is of definite interest if large tonnages exist. For example, the average grade at Endako (Canada's largest molybdenum producer) is 0.13% MoS_2.

At Trapper Lake the surface areas of known molybdenum-bearing alaskite total roughly 2.1 hectares. Much of the potentially favourable ground, such as the area between zones A and B, is covered by glacial outwash and morraine. The area of potential mineralization actually extends southward for some 2400 metres to the toe of the south glacier, and beneath it. This constitutes a large target area sufficient to contain several low grade orebodies of the minimal size envisaged to be potential producers.

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CONCLUSIONS

The alaskites of the Trapper Lake area constitute a well-defined exploration target for a large, low grade molybdenum orebody. A nine-hole drilling program totalling 1800 metres is recommended in order to cover the favourable zones on a reconnaissance basis. If significant mineralization is discovered in this work, a follow-up program of closer-spaced holes would be required. In conjunction with the drilling, prospecting and geological mapping should be carried out. New information arising from this work would be expected to influence the locations of holes remaining to be drilled. Proposed hole locations are shown on Figure 5.

RECOMMENDATIONS

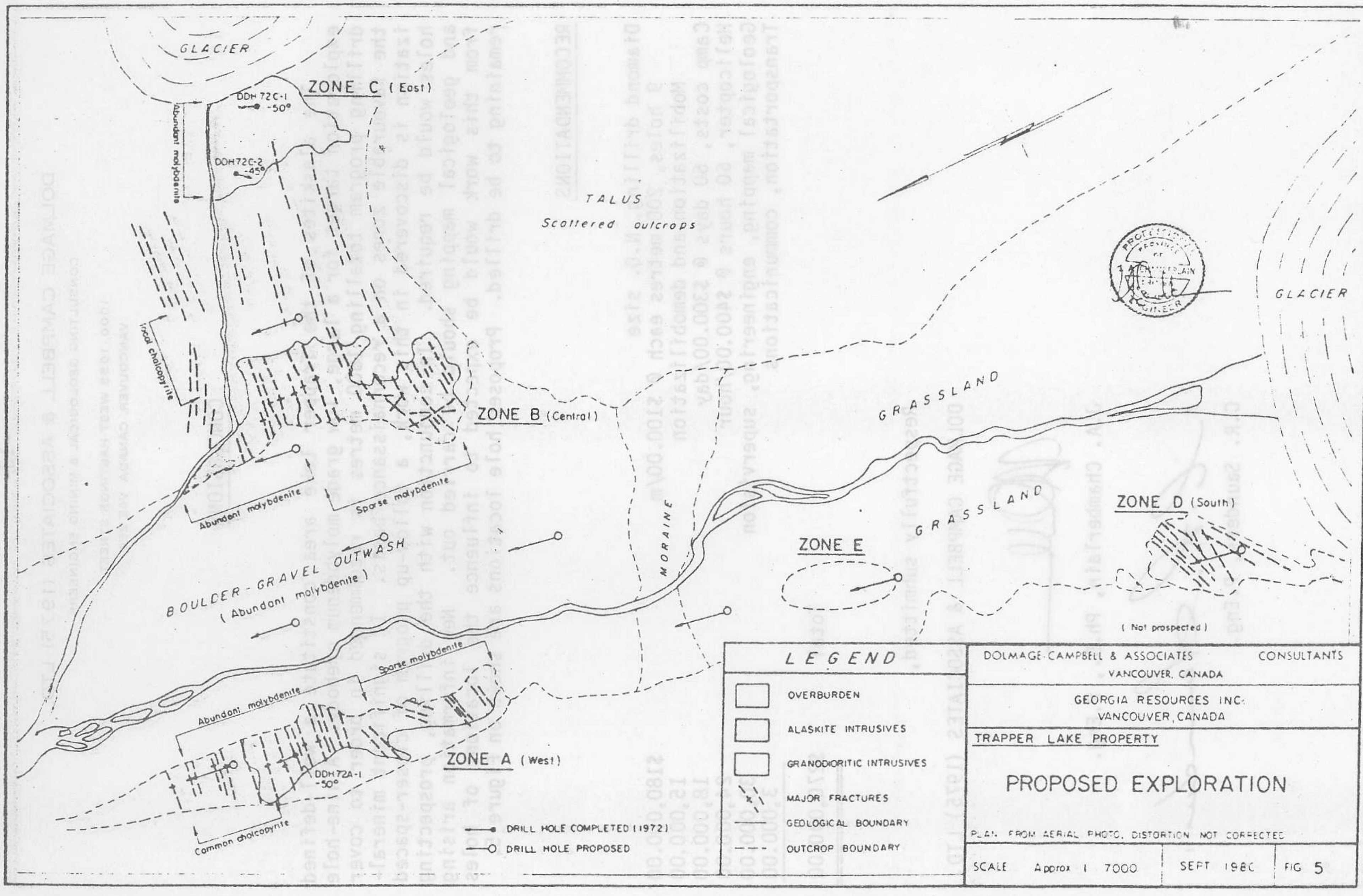
Diamond drilling, N.Q. size	
9 holes, 200 metres each @ $100.00/m	$180,000.00
Mobilization and demobilization	15,000.00
Camp costs, 60 days @ $300.00/day	18,000.00
Helicopter, 60 hours @ $400.00/hour	24,000.00
Geological mapping, engineering, supervision	30,000.00
Transportation, communications	3,000.00
Total	$270,000.00

Respectfully submitted,

DOLMAGE CAMPBELL & ASSOCIATES (1975) LTD.

J.A. Chamberlain, Ph.D., P.Eng.

C.R. Saunders, P.Eng.


GLACIER
ZONE C (East)
DDH 72C-1 -50°
DDH 72C-2 -45°
Abundant molybdenite
TALUS
Scattered outcrops
Local chalcopyrite
ZONE B (Central)
Abundant molybdenite
Sparse molybdenite
GRASSLAND
GRASSLAND
MORAINE
BOULDER - GRAVEL OUTWASH
(Abundant molybdenite)
ZONE E
ZONE D (South)
(Not prospected)
GLACIER
Abundant molybdenite
Sparse molybdenite
ZONE A (West)
DDH 72A-1 -50°
Common chalcopyrite
DRILL HOLE COMPLETED (1972)
DRILL HOLE PROPOSED
LEGEND
OVERBURDEN
ALASKITE INTRUSIVES
GRANODIORITIC INTRUSIVES
MAJOR FRACTURES
GEOLOGICAL BOUNDARY
OUTCROP BOUNDARY
DOLMAGE-CAMPBELL & ASSOCIATES CONSULTANTS
VANCOUVER, CANADA
GEORGIA RESOURCES INC.
VANCOUVER, CANADA
TRAPPER LAKE PROPERTY
PROPOSED EXPLORATION
PLAN FROM AERIAL PHOTO, DISTORTION NOT CORRECTED
SCALE Approx 1 7000
SEPT 1980
FIG 5

DOLMAGE CAMPBELL & ASSOCIATES (1975) LTD.
CONSULTING GEOLOGICAL & MINING ENGINEERS
1000-1055 WEST HASTINGS STREET
VANCOUVER, CANADA V6E 2E9

September 25, 1980.

CERTIFICATE

I, Joseph A. Chamberlain of Duncan, B.C., do hereby certify that:

1. I am a consulting geological engineer.

2. I am a graduate of the University of British Columbia, (1955), and of Harvard University, (M.A., Ph.D., in Structural and Economic Geology, 1957, 1958).

3. I am a registered Professional Engineer of the province of British Columbia.

4. From 1952 until the present, I have been engaged in regional geological studies, mining and mining exploration, engineering geology, and geological research for various companies and government institutions. I was Geologist and Research Scientist for the Geological Survey of Canada for nine years specializing on the geology of nickel, copper and uranium.

5. This report is based on examinations of assay data and geological reports listed under "References" herein.

6. I have not received, nor do I expect to receive, any interest, directly or indirectly in the Georgia Resources Inc. property described herein.

Respectfully submitted,

J.A. Chamberlain, Ph.D., P.Eng.

Vancouver,
Canada

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September 25, 1980

CERTIFICATE

I, C.R. Saunders of 666 St. Ives Crescent, North Vancouver, Canada, do hereby certify that:

1. I am a consulting geological engineer.

2. I am a graduate of the University of British Columbia, (B.A.Sc., Geological Engineering, 1956).

3. I am a registered Professional Engineer of the Province of British Columbia.

4. From 1956 until 1967 I was engaged in mining and mining exploration in Canada for a number of companies. I was Chief Geologist for Western Mines Ltd., when I left in 1967 to begin practice as a consulting geological engineer.

5. This report is based on the results of a personal examination of the Trapper Lake property made in July, 1972, and on the use of all available government and private reports, maps and records.

6. I have not received, nor do I expect to receive, any interest directly or indirectly in the property of Georgia Resources Inc. or any affiliate, nor do I beneficially own any securities directly or indirectly in Georgia Resources Inc. or any affiliate.

Respectfully submitted,

C.R. Saunders

C. Raymond Saunders, P.Eng.

Vancouver,
Canada