LEGEND

DEGERD
PLEISTOCENE AND RECENT Glacial deposits and recent alluvium; till, gravel, sand, silt, and
clay; few if any bedrock exposures
PLEISTOCENE AND/OR EARLIER
Olivine basalt; cinder cones, blocky flows, breccia, and agglomerate
MIOCENE OR PLIOCENE
9 Flat-lying olivine basalt flows; minor breccia and gravel
TERTIARY (?)
8 Conglomerate
JURASSIC AND/OR CRETACEOUS AND (?) EARLIER 7a, biotite granodiorite and granite; 7b, hornblende diorite;
7 7c, muscovite granite; 7d, biotite-hornblende syenite, biotite granodiorite, hornblende diorite, and felsite; includes septa and
inclusions of intruded rocks
6 Serpentinite
PERMIAN OR EARLIER
Greenstone, greenschist, chlorite schist, phyllite, limestone, quartz-sericite schist, quartzite, volcanic agglomerate
4a, dark grey and brown phyllite (commonly limy), limestone, sericitic quartzite; minor greenstone, quartz-feldspar-chlorite gneiss,
and meta-conglomerate; 4b, trachytic tuff and breccia
Grey and buff weathering, white, grey, and buff marble and limestone; minor greenstone and phyllite
Undivided; includes rock types common to 4a and 5; minor quartz-mica
AGE UNCERTAIN
SHUSWAP METAMORPHIC COMPLEX la, characterized by well foliated granitic gneiss; quartz-feldspar-
biotite gneiss, quartz-feldspar-hornblende gneiss, amphibolite; minor quartz-mica schist, quartzite, marble, and skarn; abundant
and locally dominant pegmatite, muscovite granite, and biotite granodiorite; lb, exclusively or dominantly biotite granodiorite;
lc, characterized by quartz-mica schistose gneiss (commonly
garnetiferous), amphibolite, quartzite, marble, and skarn; pegmatite, muscovite granite, biotite granodiorite; minor granitic gneiss;
ld, similar to unit le with abundant and locally dominant dykes and sills of pegmatite, muscovite granite, and biotite granodiorite;
le, undivided, may include all rock types found in units la and lc. The granitic rocks may be equivalent to those of 7
Small rock outcrop
Geological boundary (defined, approximate, and assumed)
Foliation including rock classage, achietosity, maissagity
Foliation including rock cleavage, schistosity, gneissosity, and bedding (inclined, vertical, and horizontal)
Foliation including rock cleavage, schistosity, gneissosity, and bedding (inclined, vertical, and horizontal)
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(SEYMOUR ARM, WEST HALF)

BRITISH COLUMBIA

Scale 1:253,440 (1 inch to 4 miles)

that further prospecting, using modern techniques of geochemistry, may be warranted. Jones, A.G.: Vernon Map-area British Columbia; Geol. Surv. Can., Mem. 296. ²Campbell, R. B.: Quesnel Lake East Half, British Columbia; Geol. Surv. Can.,

³Lowdon, J.A., et al.: Age determinations and geological studies; Geol. Surv. Can.,

Paper 62-17, p. 7.

Wheeler, J.O.: Big Bend, British Columbia; Geol. Surv. Can., Preliminary Report

82M010-06 MAP 48 1963

ADAMS LAKE RITISH COLUMBIA 82 M W

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of Vernon map-area1. The complex has been divided into sub-units as indicated on

The Shuswap Metamorphic Complex (1) is equivalent to the Monashee Group

DESCRIPTIVE NOTES

and Barriere Rivers and of Sinnax Creek. Logging and Forest Service roads provide access to the valleys of Raft, Mid, and Adams Rivers, to both branches of Scotch Creek, to several smaller vallers, and onto many hillsides. Much of the map-area is covered by heavy timber and lense underbrush and in general travel is difficult

and bedrock exposures are meatre. Forest growth extends to an elevation of about

The main public roads in the map-area lie in the valleys of North Thompson

the map, but clearly definable boundaries between them are lacking and none is shown on the map. In zones between sub-unit symbols on the map the rock assemblages change gradationally from that of one sub-unit to that of another. No direct evidence regarding the age of the rocks or of the metamorphism of the complex was found. The rocks seem certainly to be older than Carboniferous and they may be at least partly equivalent to the Proterozoic Kaza Group and lower Palaeozoic Cariboo Group, which trend toward the map-area from the northwest². The rocks of the complex may have been deformed and possibly metamorphosed

prior to the deposition of those of units 2 to 5 but probably the most intense deforma-

tion and metamorphism took place in post-Permian time³.

The rocks of the Shuswap Metamorphic Complex (1) are a strongly foliated and lineated assemblage of metasedimentary gneisses and schists intruded by an enormous number of dykes, sill, and small irregular bodies of granitic rocks. Granitic rocks included in unit are not grouped with those of unit 7, with which they may be equivalent, because they cannot be mapped separately and shown independently at the present scale. Pegmatite comprises more than 70 per cent of the exposed rocks in some places, for example on the slopes of Adams River valley south of Tumtum Lake. Huscovite granite and aplite appear to be closely related to pegmatite in origin and occurrence. Unfoliated or weakly foliated, mainly medium-grained biotite granodicrite forms large masses, as directly east of Trophy Mountain, and a multitude of dykes and sills, mainly in the northern part of the map-area.

Granitic rocks, including pegmatite, can be seen to be intrusive into dilational openings in many places. Where the sequence of intrusion could be determined it was found that permatite clearly intrudes all other granitic rocks.

The rocks of units 2 to 5 are equivalent to the Eagle Bay Formation of the Mount Ida Group¹ and to the Adams Lake and Nisconlith Series of Dawson⁴. The map-units cannot be regarded as distinct stratigraphic units nor are their relative ages necessarily indicated by their position in the legend. The limestone bands of unit 3 are not likely to represent a single stratigraphic unit and it may be that no one band is in the same stratigraphic position as any other.

Fossils from a limestone of unit 3 are presumably of Carboniferous or Permian age, and that, in the writer's opinion, may be the age of all the rocks included in units 2 to 5.

The contact of the rocks of units 2 to 5 with those of the Shuswap Metamorphic Complex was not directly observed. In the southern part of the map-area the contact appears to lie along system of north-south and east-west trending faults whereas in the northern part it may be a combination of faults and an unconformity, though evidence of the latter is lacking.

In the southern part of the map-area the rocks of unit 4a are mainly well-foliated and strongly crenulated argillaceous phyllites with associated thin beds of limestone. To the north quartzose rocks are more prevalent. Quartzfeldspar-chlorite gneiss is exposed on the west side of Scotch Creek south of the forks and is believed to be the metamorphic derivative of quartz-pebble conglom-

Metavolcanic rocks are characteristic of, but not necessarily dominant in, the rocks of unit 5. Much of the unit is composed of green phyllite and chlorite schist which may be all or partly of sedimentary origin. Other rocks of the unit are distinctly meta-sedimentary. Foliation is strongly developed in these rocks but lineations are confined to the phyllite and schist and are apparently lacking in greenstone.

Unit 2 cannot be easily characterized by any particular rock type; rather it includes types typical of both units 4a and 5, though they are dominantly metasedimentary. It may be that unit 2 represents the rocks of units 4a and 5 folded together in a distinctive pattern. With better exposure it might be possible to divide these rocks into

lithologic units equivalent to units 4a and 5.

Serpentinite (6) associa ed with greenstone (5) was found to contain very short and sparse asbestos fibre in one place.

Many dykes of granodiorite, granite, and rhyolite, too small to show on the map are found on Adams Plateau and are evidently related to the nearby small granitic The assorted rocks of unit 7d are intruded in a long narrow belt which may be

the locus of a zone of fracturing. The intrusions are evidently in the form of dykes between which are septa of the invaded rocks. All the metamorphic rocks of the map-area are cut by basaltic dykes from a

few inches to several feet wide. These dykes characteristically trend north or a few degrees east of north. They are perhaps, most common in the eastern part of the map-area. Well-consolidated congruence (8) evidently formed sometime during the

Tertiary from granitic detritus derived from hills to the west of Adams Lake valley.

The basaltic flows of unit 9 are equivalent to valley flows along Clearwater

River west of the map-area, where they underlie lava and agglomerate of unit 10. The latter are associated with cinder cones, and though these deposits have been glaciated, similar cones and flows of post-lacial age were noted nearby in Quesnel Lake map-

All the metamorphic rooks are intensely deformed and those of the Shuswap Metamorphic Complex (1) have been found by Wheeler⁵ to have suffered at least two periods of folding; first into recumbent, attenuated folds on axes trending a little north of east, and second into more open and upright folds on northwesterly axes. Evidence of two periods of folding is found only locally in the rocks of units 2 to 5. In the writer's opinion these rocks are isoclinally folded, and the axial planes of the folds are roughly parallel with the foliation. The axes of such folds apparently trended somewhat north of west. Subsequently the strata may have been redeformed into broad warps and local minor folds.

The majority of mineral deposits in the map-area are confined to the rocks of units 2 to 5. Not all the known deposits are shown on the map but those that are indicate the general areas where the known deposits are concentrated. No deposit in the area has, as yet, proved to be of economic value, with the possible exception of placer gold on Scotch Creek. The concentration of mineral deposits in the area extending northwesterly from Adams Plateau plus the limited bedrock exposure is a combination that suggests

Map 1-1963.

Dawson, G. M.: Shuswap Sheet, British Columbia; Geol. Surv. Can., Map 604.

in preparation.